



Does the hook structure constrain future precipitation extremes and floods?

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In collaboration with Prof. Pierre Gentine, Prof. Louise Slater and Prof. Shenglian Guo



1 Background

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Save reefs to rescue all ecosystems

An approach that tackles the underlying causes of coral-reef decline could be applied to other habitats, argue Tiffany H. Morrison, Terry P. Hughes and colleagues.

Nature Alert: All the coral reefs in the world might be gone by 2070 if global heating continues on its current path.



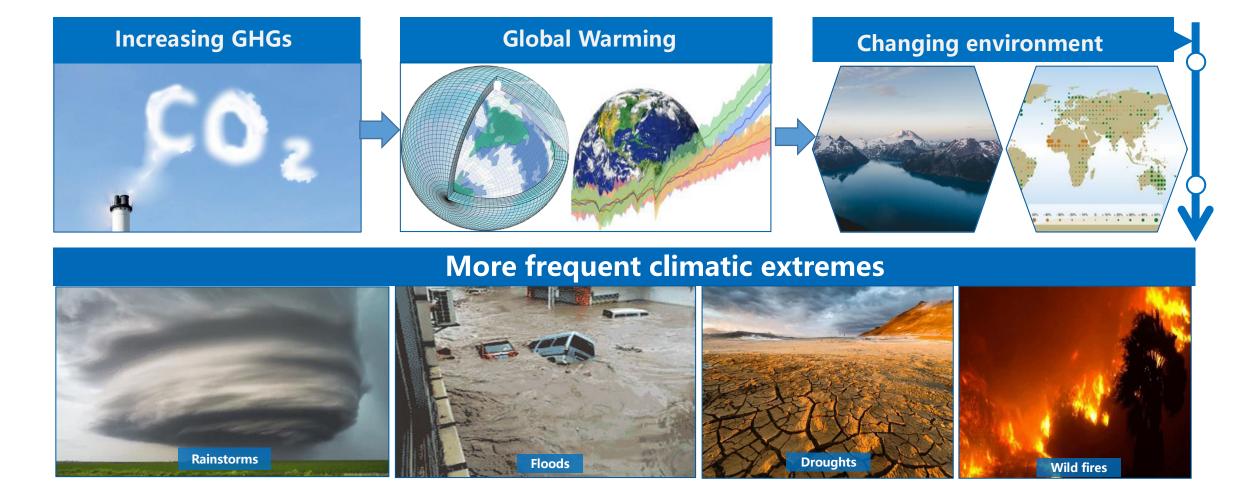
Climate Warming

2019-2020: Australian wildfires have killed half a billion animals and plants.

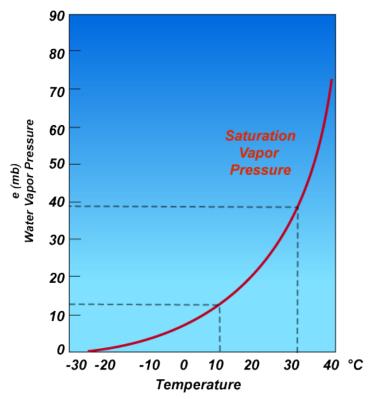


1 Background

Anthropogenic warming due to increasing greenhouse gas emission has altered the climate system and water-carbon cycle



Atmospheric water vapor holding capacity should increase with rising temperature



$$e_{\rm S}(T) = e_{\rm S0} \exp(\frac{L_v}{R_v} \left[\frac{1}{T_0} - \frac{1}{T} \right])$$
 saturated vapor pressure

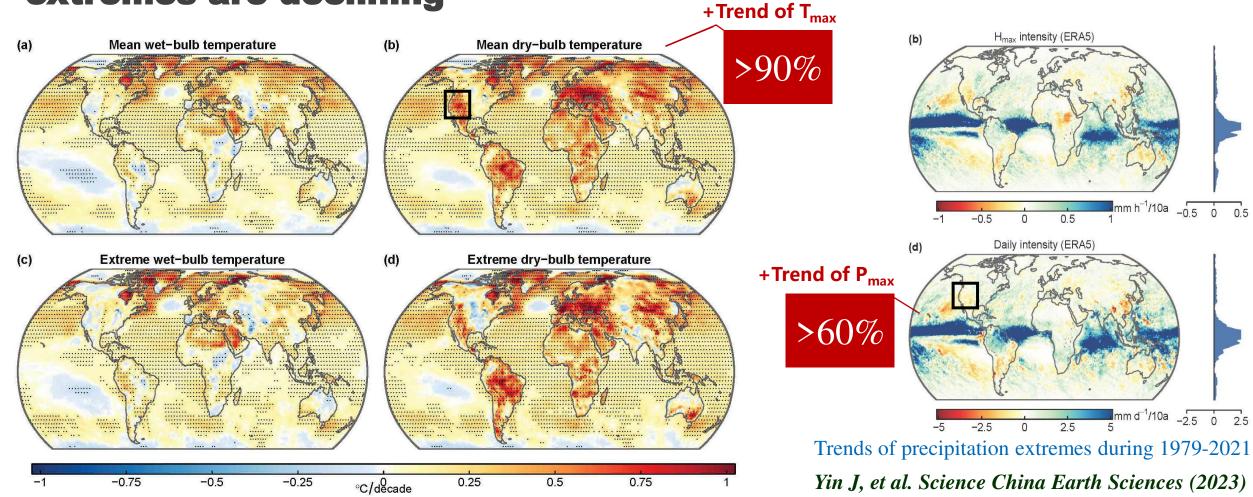
$$\frac{de_S}{e_S dT} = \frac{L_v}{R_v T^2}$$

Latent heat of vaporization

Clausius-Clapeyron

- (1) saturation vapor pressure is dependent solely on temperature
- (2) 6.8 %/K at 25 °C (usually called as C-C scaling)

In many regions, temperature is rising, while precipitation extremes are declining

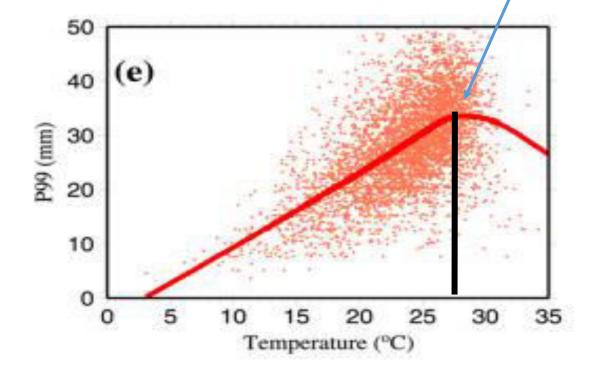


Global trend of mean and extreme temperatures during 1979-2021

Yin J, et al. GRL (2022)

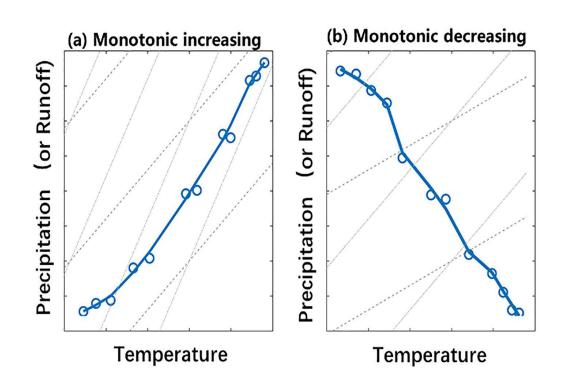
Unlike the thermodynamic expectations, we find a hook structure between precipitation/runoff extremes and temperatures

Hook structure: Extremes strengthen with rising temperature up to a peak point (T_{pp}) and decline thereafter

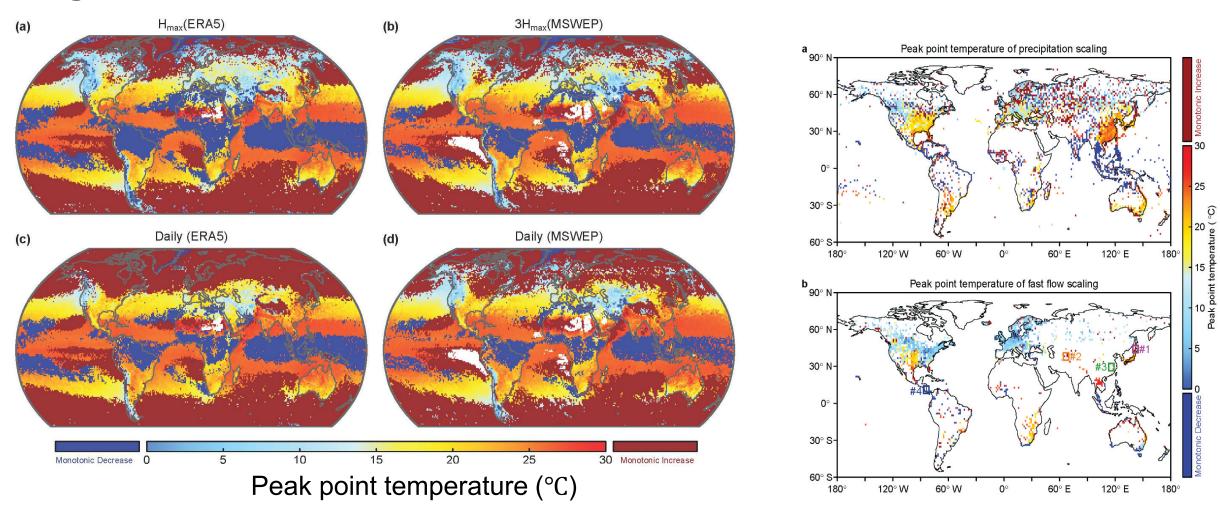


Binning Scaling

$$EPS = (\exp^{\frac{\ln P_b - \ln P_a}{T_{b-}T_a}} - 1) \times 100\%$$



The underlying causes of this widely reported hook structure is not yet well-understood



Yin J, et al. Science China Earth Sciences (2023)

Yin J, et al. Nature Communications (2018)

The underlying causes of this widely reported hook structure is not yet well-understood

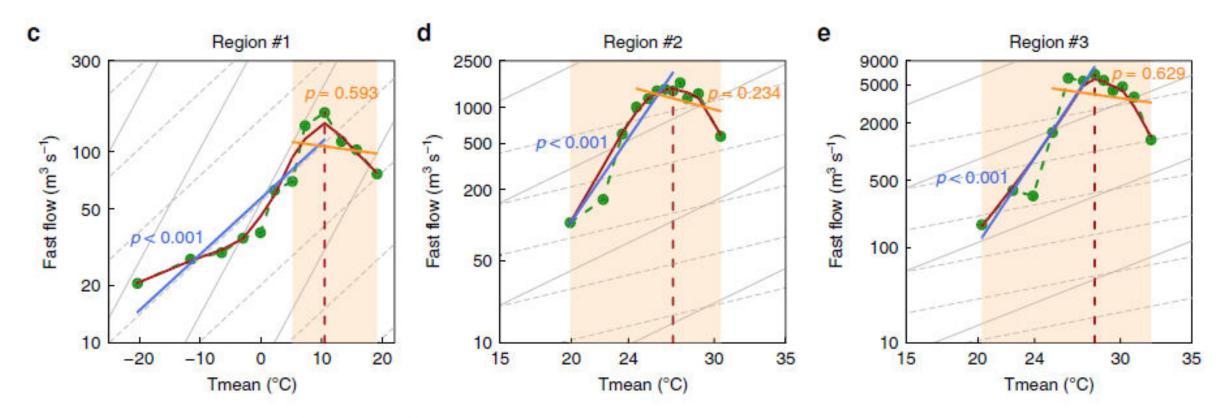


Figure: Temperature scaling of storm runoff over three typical catchments

Why can we observe a hook structure?

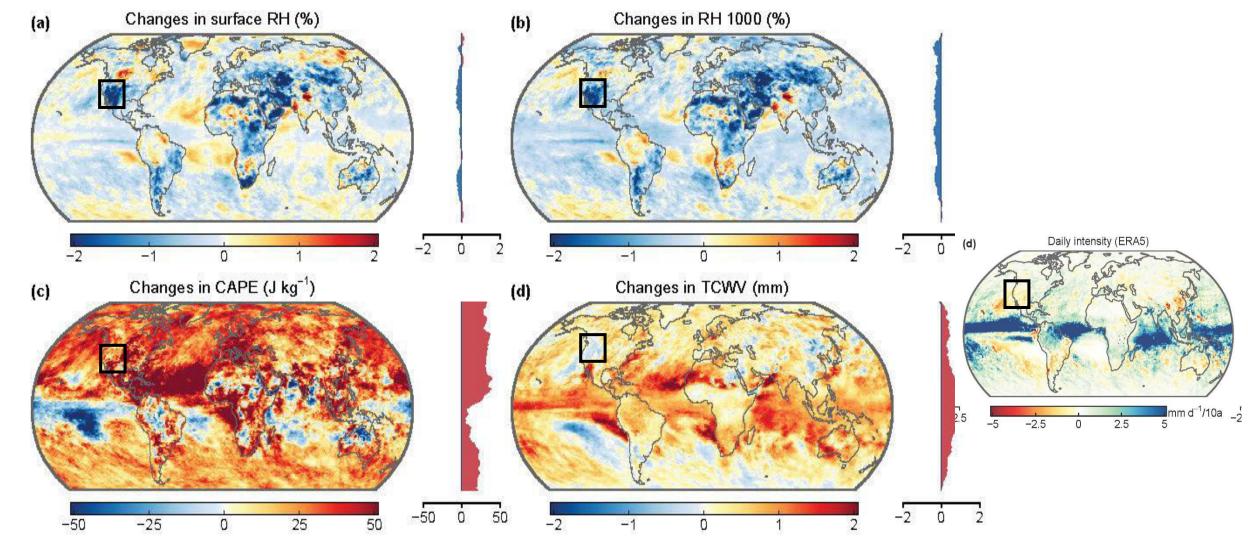


Figure: Trends in relative humidity, CAPE, and Total column water vapor (TCWV) during precipitation extremes

Why can we observe a hook structure?

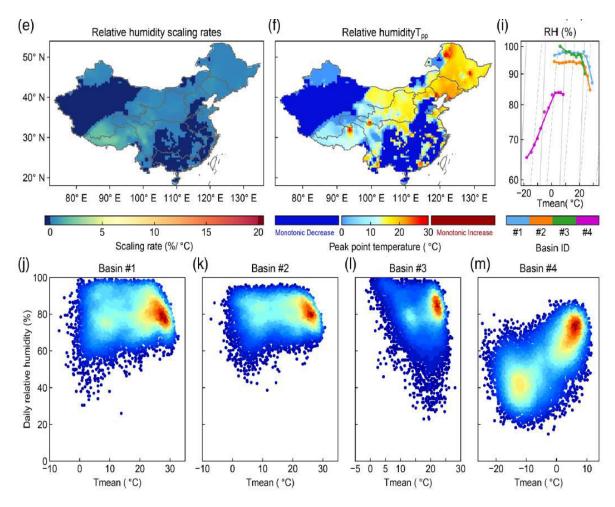


Figure: Scaling pattern between Relative humidity and temperatures

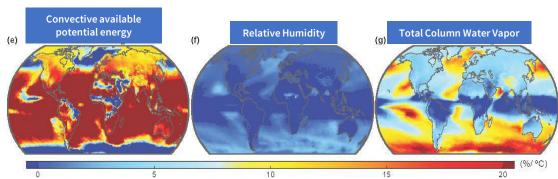


Figure: Temperature scaling of CAPE, RH and TCWV

■ We find negative (sub C-C) scaling rates of *Relative Humidity* and *Total Column Water Vapor (TCWV)*, implying continental moisture limitations in hotter environments. Atmospheric dynamics might constrain extremes intensification.

Yin J, et al. Water Resources Research (2021)

Why can we observe a hook structure?

Divergent sensitivities of P_e are reported

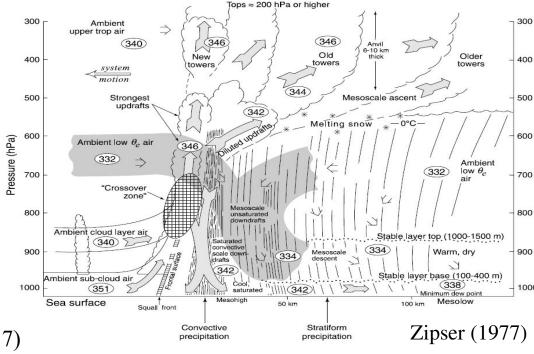
C-C scaling only tells part of the story

Local vertical motion

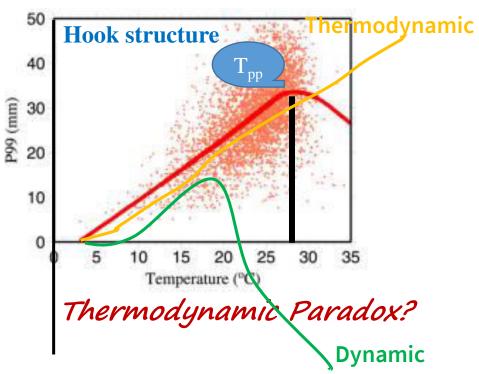
Available atmospheric moisture

$$P_e \sim -\left\{\omega_e \frac{dq_s}{dp}\bigg|_{\theta^*}\right\}$$
 O'Gorman et al. (2012, 2017)

Precipitation involves complicated systems and coupled processes. How these physical processes impact Extreme precipitation sensitivity (*EPS*)?

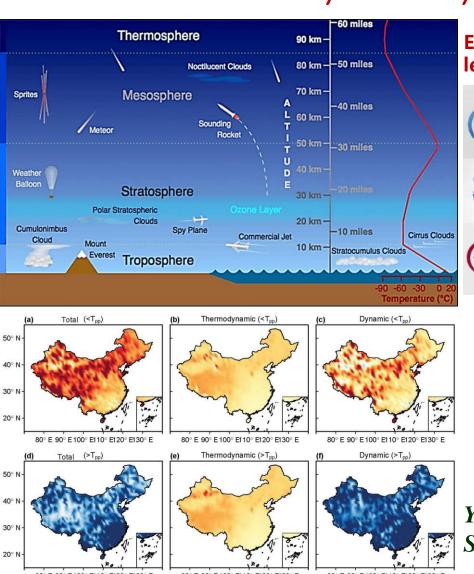


■ The physical-based diagnose model is used to quantify the thermodynamic and dynamic contributions of EPS, which is based on *energy budged balance* during precipitation processes.



$\frac{\delta P_{\rm e}}{P_{\rm e}} \approx \frac{\delta \{\omega(\partial q_{\rm sat}/\partial p)\}}{\{\omega(\partial q_{\rm sat}/\partial p)\}}$ $\frac{\delta P_{\rm e}}{P_{\rm e}} \approx \frac{\{\omega\delta(\partial q_{\rm sat}/\partial p)\}}{\{\omega(\partial q_{\rm sat}/\partial p)\}} + \frac{\{\delta(\omega)(\partial q_{\rm sat}/\partial p)\}}{\{\omega(\partial q_{\rm sat}/\partial p)\}}$ $\frac{\delta P_{\rm e}}{P_{\rm e}} \approx \frac{\{\omega\delta(\partial q_{\rm sat}/\partial p)\}}{\{\omega(\partial q_{\rm sat}/\partial p)\}} + \frac{\{\omega(\partial q_{\rm sat}/\partial p)\}}{\{\omega(\partial q_{\rm sat}/\partial p)\}}$ $\frac{\delta P_{\rm e}}{P_{\rm e}} \approx \frac{\{\omega\delta(\partial q_{\rm sat}/\partial p)\}}{\{\omega(\partial q_{\rm sat}/\partial p)\}} + \frac{\{\omega(\partial q_{\rm sat}/\partial p)\}}{\{\omega(\partial q_{\rm sat}/\partial p)\}}$ $\frac{\delta P_{\rm e}}{P_{\rm e}} \approx \frac{\{\omega\delta(\partial q_{\rm sat}/\partial p)\}}{\{\omega(\partial q_{\rm sat}/\partial p)\}} + \frac{\{\omega(\partial q_{\rm sat}/\partial p)\}}{\{\omega(\partial q_{\rm sat}/\partial p)\}}$ $\frac{\delta P_{\rm e}}{P_{\rm e}} \approx \frac{\{\omega\delta(\partial q_{\rm sat}/\partial p)\}}{\{\omega(\partial q_{\rm sat}/\partial p)\}} + \frac{\{\omega(\partial q_{\rm sat}/\partial p)\}}{\{\omega(\partial q_{\rm sat}/\partial p)\}}$

We focus on moisture transport in Troposphere



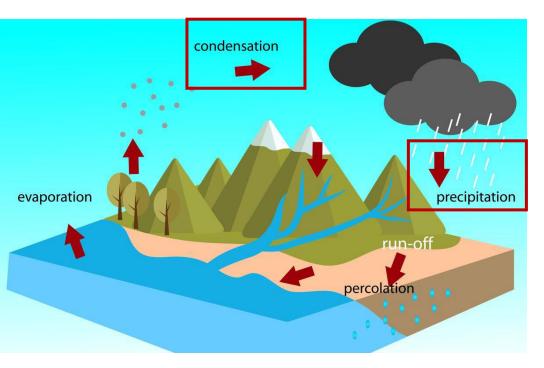
ERA5 data pressure levels (1-1000hPa)

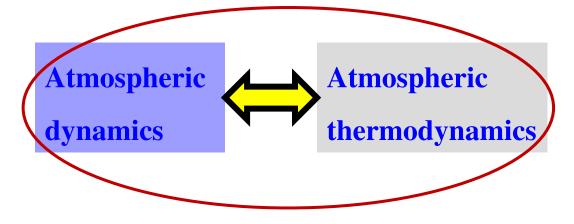


Yin J, et al. Chinese Science Bulletin (2021)

Thermodynamic contributes to precipitation intensification, while atmospheric dynamics decline precipitation in hotter

environment





Physical-diagnostic model

$$\frac{\delta P_{\rm e}}{P_{\rm e}} \approx \frac{\{\omega\delta(\partial q_{\rm sat}/\partial p)\}}{\{\omega(\partial q_{\rm sat}/\partial p)\}} + \frac{\{\delta(\omega)(\partial q_{\rm sat}/\partial p)\}}{\{\omega(\partial q_{\rm sat}/\partial p)\}}$$
Thermodynamic Dynamic

Do all the thermodynamic components always intensify precipitation extremes?

Do all the thermodynamic components always intensify precipitation extremes? Can we know more details?

By further decomposing moist-adiabatic derivative of saturation specific humidity

$$P_e \sim -\left\{\omega_e \frac{dq_s}{dp}\Big|_{\theta^*}\right\}$$

$$P_{the} \sim -\left\{\omega_{avg} \frac{dq_s}{dp}\bigg|_{\theta^*}\right\} = -\left\{\omega_{avg} \left(\frac{\partial q_s}{\partial p} + \frac{\partial q_s}{\partial T} \cdot \frac{dT}{dp}\bigg|_{\theta^*}\right)\right\} = -\left\{\omega_{avg} \frac{\partial q_s}{\partial p}\right\} - \left\{\omega_{avg} \frac{\partial q_s}{\partial T} \cdot \frac{dT}{dp}\bigg|_{\theta^*}\right\}$$



Pressure component

$$pPR = -\left\{\omega_{avg} \frac{\partial q_s}{\partial p}\right\}$$

$$pPR = -\left\{\omega_{avg} \frac{\partial q_s}{\partial p}\right\} \qquad \begin{array}{c} \textbf{Temperature} \\ \textbf{component} \end{array} \qquad pT = -\left\{\omega_{avg} \frac{\partial q_s}{\partial T} \cdot (\frac{dT}{dp}\Big|_{\theta^*})_{avg}\right\}$$

Lapse Rate component

$$pLR = -\left\{ \omega_{avg} \left. \frac{\partial q_s}{\partial T} \cdot \frac{dT}{dp} \right|_{\theta^*} \right\} - \left(-\left\{ \omega_{avg} \left. \frac{\partial q_s}{\partial T} \cdot \left(\frac{dT}{dp} \right|_{\theta^*} \right)_{avg} \right\} \right)$$

Gu, Yin*, Nature Communications (2023)

Does the thermodynamic always intensify precipitation extremes? Can we know more details?

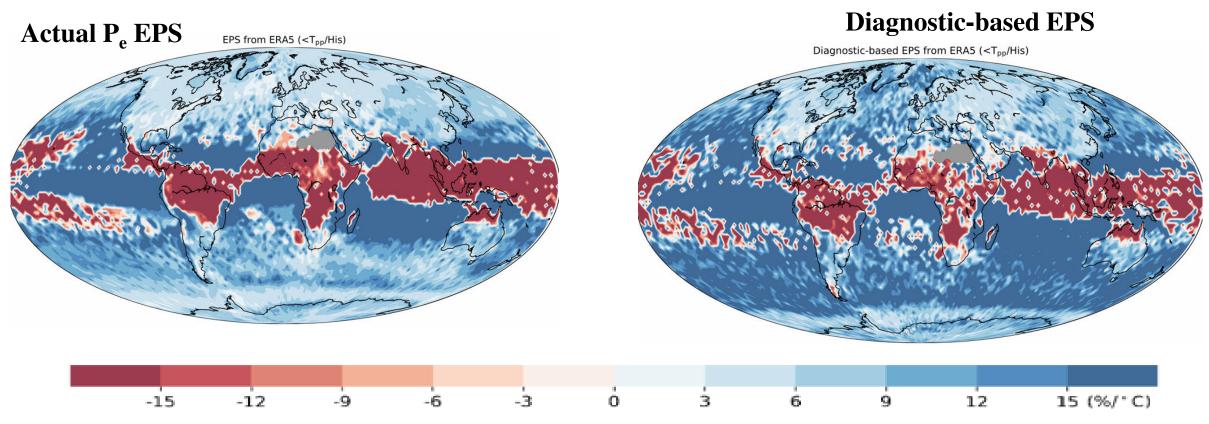


Figure: Evaluation of simulation performance of physically diagnostic model

Gu, Yin*, Nature
Communications (2023)

The thermodynamic components do not always contribute to precipitation intensification due to changes in temperature laps rate

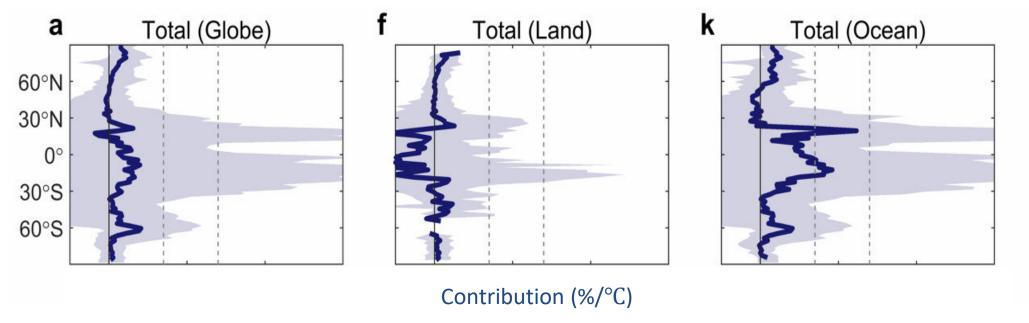
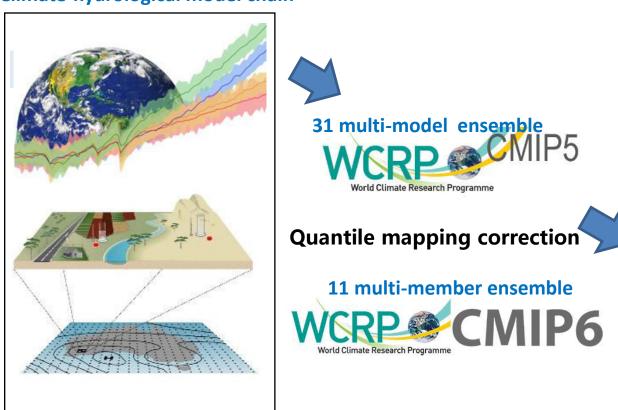
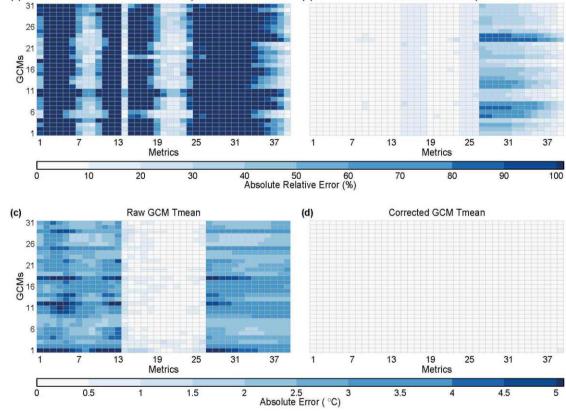


Figure: Zonal EPS anomalies between the reference and future periods

Does the hook structure imply a potential upper bound for future extremes?

Climate-hydrological model chain





Corrected GCM Prep

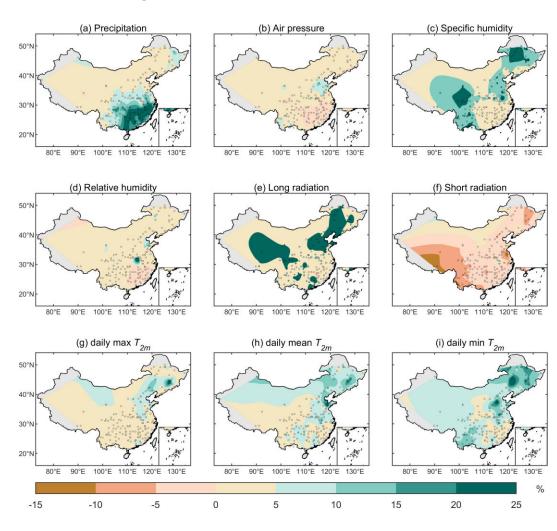
Raw GCM Prep

Yin J*, et al. Nature Sustainability (2023)

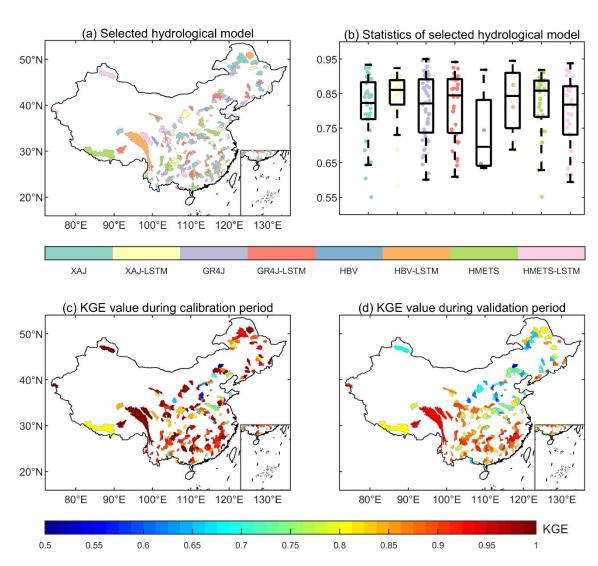
Figure: Bias correction performance of climate outputs

We propose a hybrid HM-LSTM model Yin J*, et al. Nature Sustainability (2023) Horizontal Grid (Latitude-Longitude Vertical Grid RP (SSP1-26) **Global climate model outputs Downscaling methods Climate** Runoff variables simulations RF algorithms /projections **Sensitivity KEY Analysis Hydrological model (HM)** Flood assessment

Sensitivity of climate variables to runoff

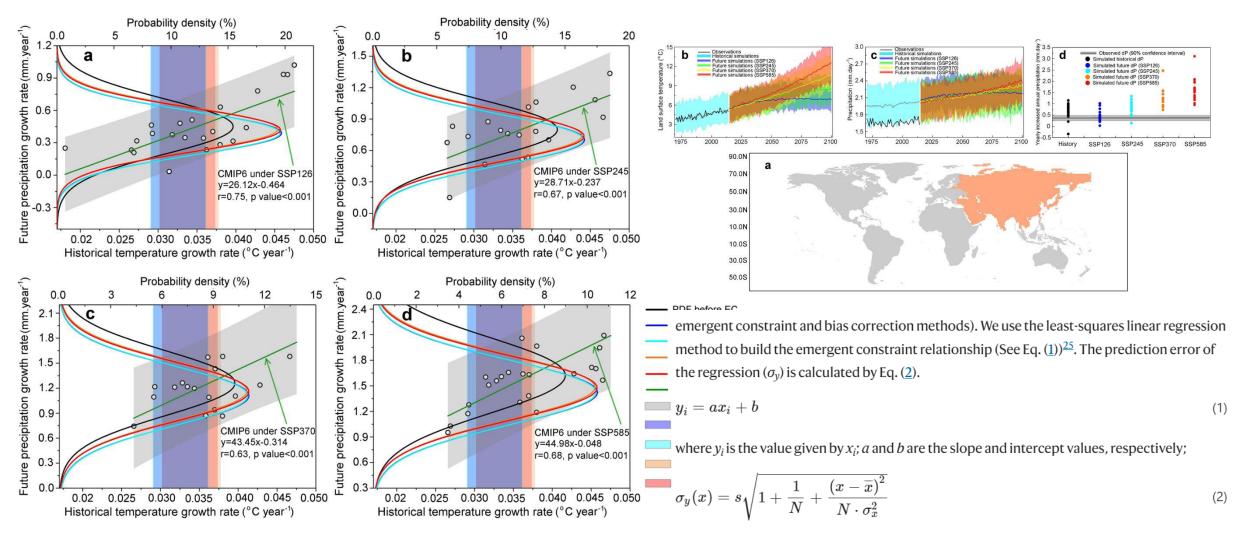


Kang S, Yin J*, et al., 2023 Earth's Future



Performance of machine Learning-constrained runoff Simulations

Future projection of precipitation, evaporation and snow cover by Emergent Constraint



Cai, Yue, Slater, Yin et al. Nature Communications (2022)

Does the hook structure imply a potential upper bound for future extremes?

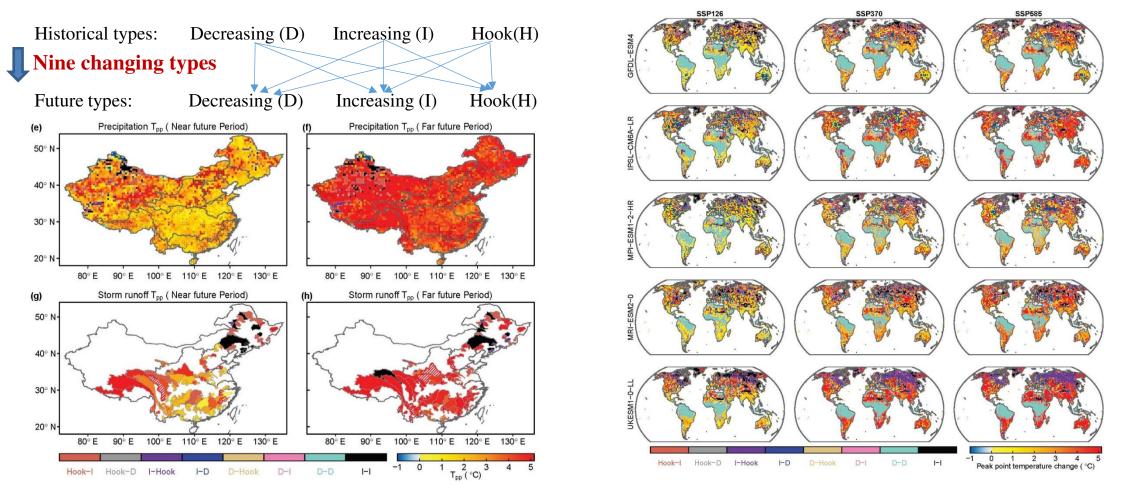


Figure: The GCM ensemble mean change of scaling structures and T_{pp} relative to historical period

Does the hook structure imply a potential upper bound for future extremes?

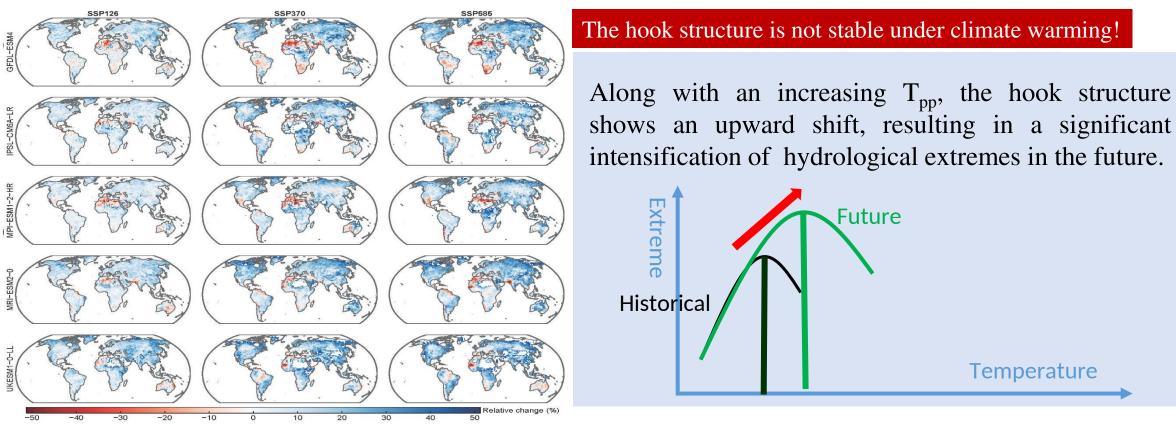
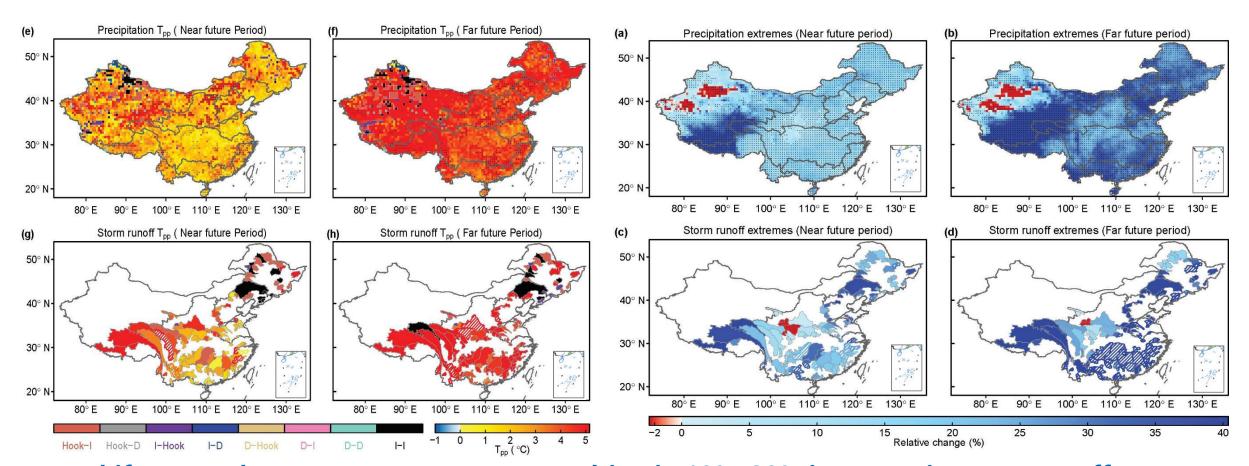


Figure: Relative changes of precipitation extremes in future climates

More than 80% of continental areas retain a hook structure during historical and future periods, while Peak point temperature (T_{pp}) progressively increases with atmospheric warming.

Does the hook structure imply a potential upper bound for future extremes?



 T_{pp} shifts toward warmer temperatures, resulting in 10%–30% increases in storm runoff extremes.

Yin J, et al., WRR (2021 Editors' Choice Awards)

Conclusions

- Precipitation and storm runoff extremes show a hook structure with temperature rising.
- Atmospheric thermodynamic intensifies precipitation extremes, while atmospheric dynamics decline precipitation in hotter environment.
- ☐ The thermodynamic components do not always contribute to precipitation intensification due to changes in temperature laps rate.
- □ The hook structure does not imply potential bound for future extreme intensification.





Thank you!

Research interests:

Atmospheric dynamics; Climate change;

Ecohydrology; Machine learning;

Hydrological extremes; Remote sensing

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