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Potential Evaporation and the Complementary Relationship

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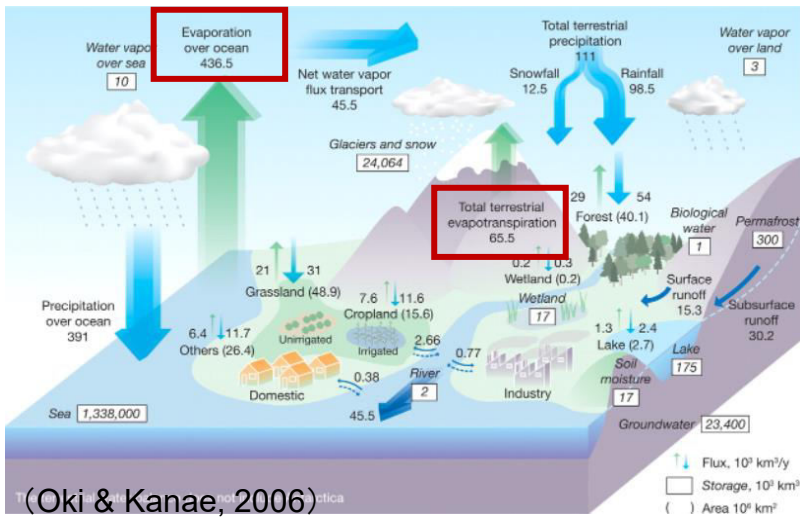
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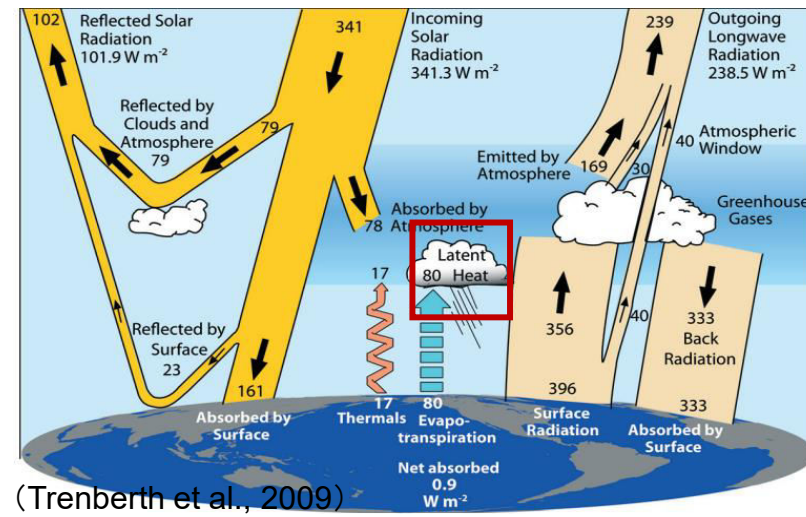
Evaporation (E) is a crucial linkage between water and energy

(Oki & Kanae, 2006; Trenberth et al., 2009)

cycles



(Oki & Kanae, 2006)



(Trenberth et al., 2009)

Water Budget $P = E + R + \Delta S$

The 2nd largest water flux in the terrestrial water cycle

Energy Budget

$R_n = LE + H + G$

Consuming 82% of the energy gained on the Earth's surface

■ Potential Evaporation is a key concept to estimate E

- Actual Evaporation :

The evaporation under natural conditions



- **Potential Evaporation :**

The evaporation occurs with unlimited water supply



Models for estimating the potential evaporation

Combination method (Penman, 1948) Taylor, 1972)

$$LE = \frac{\Delta R_n + \rho_a C_p \frac{D_a}{r_a}}{\Delta + \gamma}$$

areodynamic term + radiative term

Coupling the aerodynamic and radiative term, it is considered to be a more reliable model.

Energy-based method (Priestley &

$$LE = \alpha \frac{\Delta}{\Delta + \gamma} R_n$$

equilibrium evaporation

minimal advection

minimal areodynamic term + radiative term

Simplifying aerodynamic term with minimum advection, it has good performance in E_p estimation

Problems of classic potential evaporation

- potential evaporation:** The evaporation that would occur when the surface is saturated

Priestley-Taylor $LE = \alpha \frac{\Delta}{\Delta + \gamma} R_n$

Penman $LE = \frac{\Delta R_n + \rho_a C_p \frac{D_a}{r_a}}{\Delta + \gamma}$

Meteorological forcing variables observed in real conditions do not necessarily represent that measured over hypothetical wet surfaces



Temperature < Temperature

Solar radiation ≈ Solar

radiation

Longwave radiation > Longwave

radiation

Humidity > Humidity

Problems of classic potential evaporation (Brutsaert, 2015)

models

Potential Evaporation (E_{po})

(large) surface is wet and the surrounding air is humid



Apparent Potential Evaporation (E_{pa})

(small) surface is wet but the surrounding air is dry



Temperature < Temperature

Solar radiation ≈ Solar

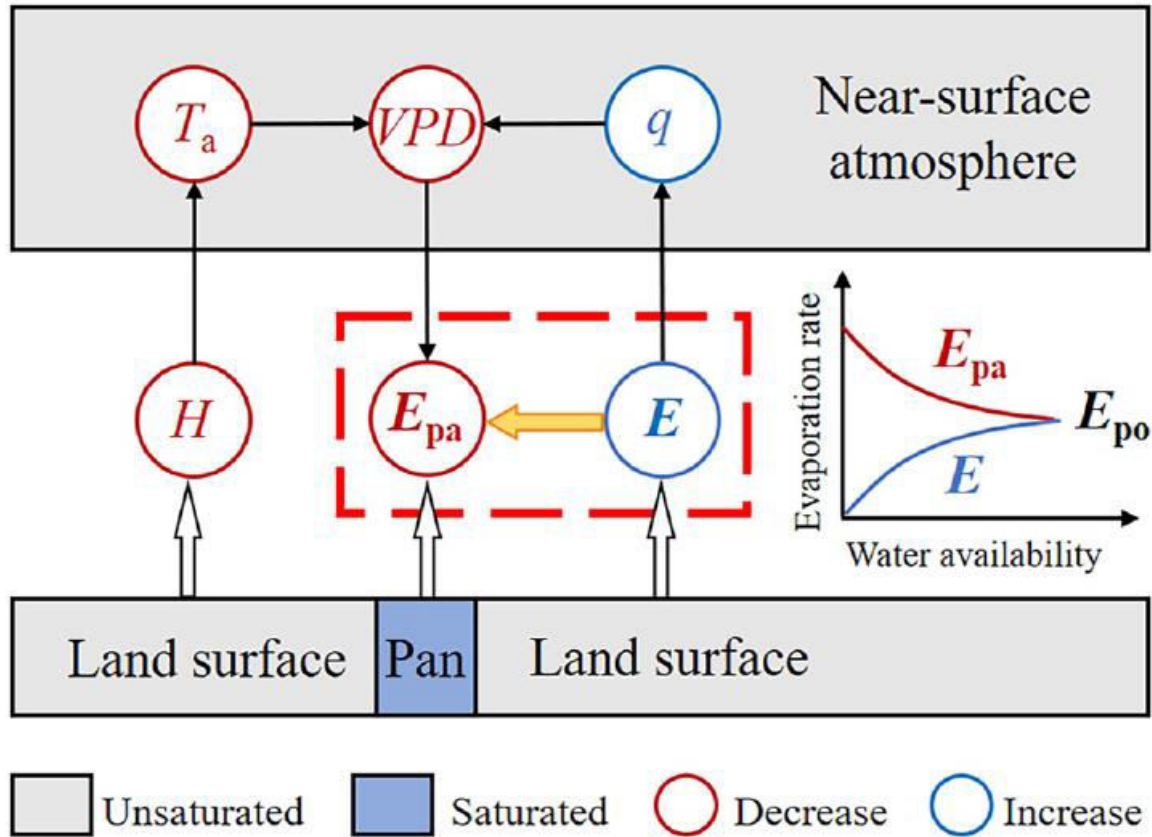
radiation

Longwave radiation > Longwave

radiation

Humidity > Humidity

■ The Complementary



(Bouchet, 1963)

The CR provides a framework for estimating E with basic meteorological observations based on:

$$E = f(E_{pa}, E_{po})$$

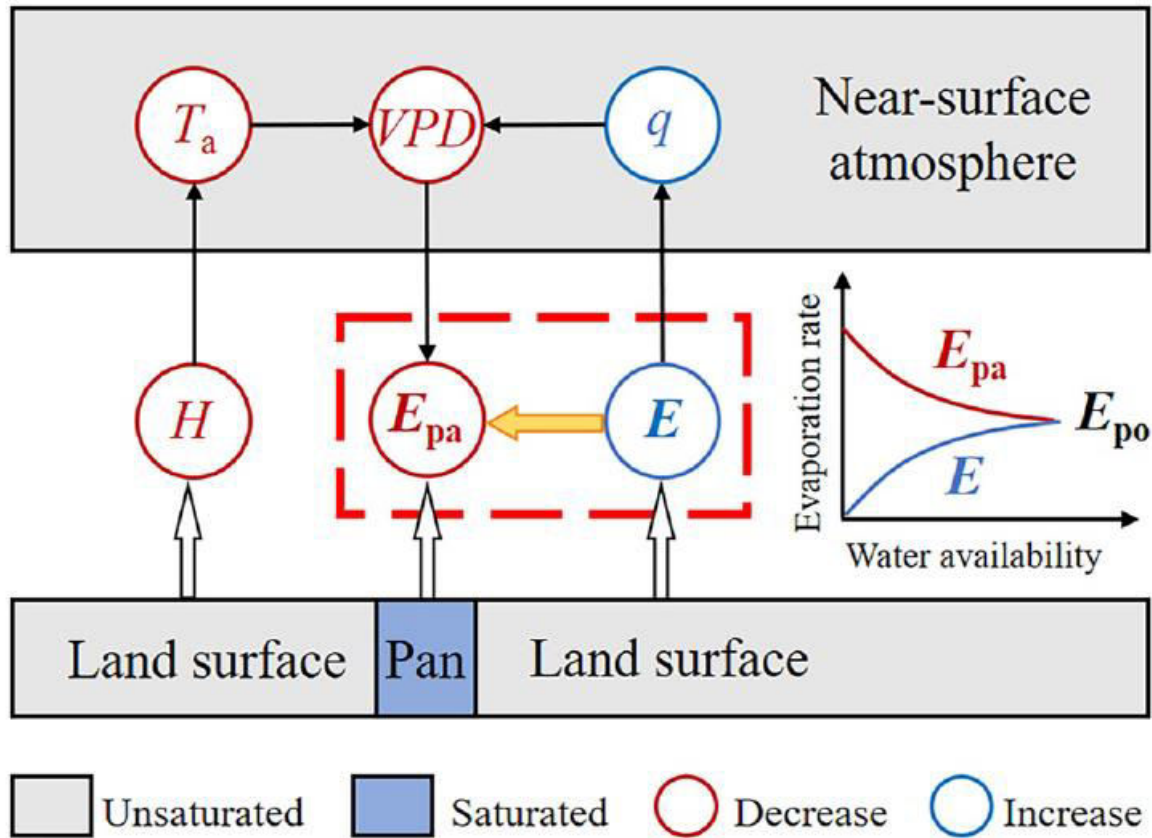
Apparent Potential Evaporation (E_{pa}):

(small) surface is wet but the surrounding air is dry

Potential Evaporation (E_{po}):

(large) surface is wet and the surrounding air is humid

■ The Complementary



(Bouchet, 1963)

Apparent Potential Evaporation (E_{pa}):

(small) surface is wet but the surrounding air is dry

Pan observation/ Penman equation

Potential Evaporation (E_{po}):

(large) surface is wet and the surrounding air is humid

Debated!

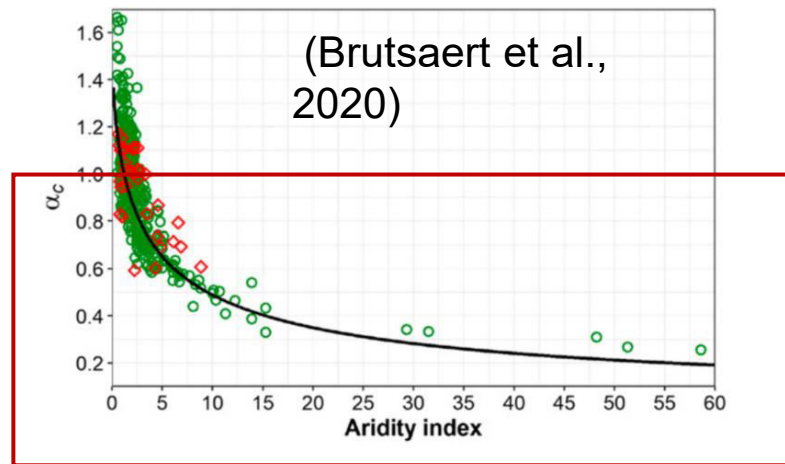
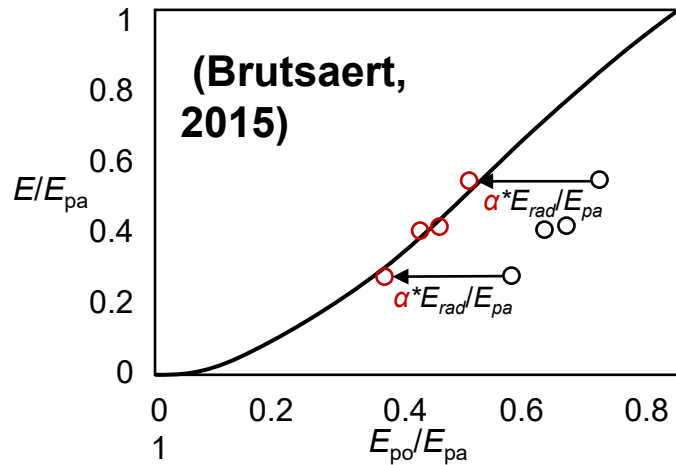
Meteorological variables observed in real conditions do not necessarily represent that measured over wet surfaces

■ Problems of E_{po} estimation in the CR

• Polynomial CR (Brutsaert, 2015)

$$\frac{E}{E_{pa}} = 2\left(\frac{\alpha E_{rad}}{E_{pa}}\right)^2 - \left(\frac{\alpha E_{rad}}{E_{pa}}\right)^3$$

$$E_{po} = \alpha \frac{\Delta}{\Delta + \gamma} (R_n - G) = \boxed{\alpha}^* E_{rad}$$



Problem: (Phillip, 1987)

**Fitted α lower than 1
 α should be higher**

**than 1 internal
inconsistency!**

Problems of E_{po} estimation in

the CR

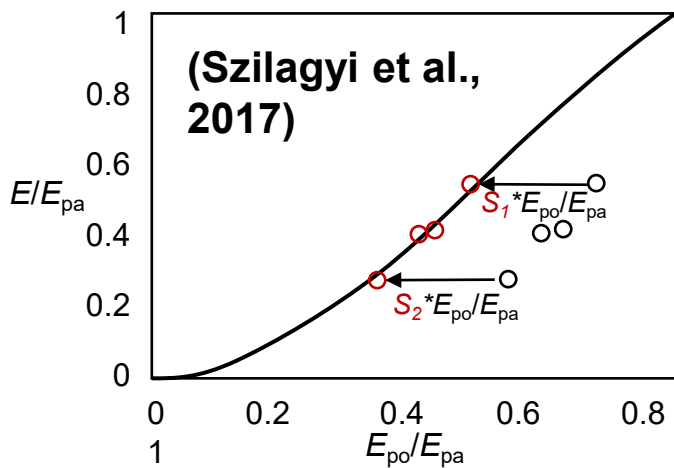
- Rescaled CR (Szilagyi et al., 2017) $E_{po} = \alpha \frac{\Delta(T_{ws})}{\Delta(T_{ws}) + \gamma} (R_n - G)$ **Derive the wet surface temperature**

$$\frac{E}{E_{pa}} = 2\left(S \frac{E_{po}}{E_{pa}}\right)^2 - \left(S \frac{E_{po}}{E_{pa}}\right)^3$$

$$S = \frac{E_{pa_max} - E_{pa}}{E_{pa_max} - E_{po}}$$

Assume a constant R_n

$$\beta = \frac{H}{LE} \approx \frac{R_n - G - LE_{P_PO}}{LE_{P_PO}} \approx \gamma \frac{T_{ws} - T_a}{e_s(T_{ws}) - e(T_a)}$$



Problem:

Ignore the interdependence between R_n

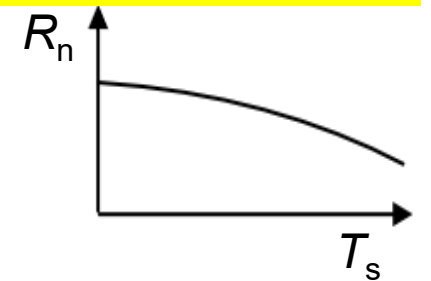
Energy Budget :

$$R_n = R_{si} - R_{so} + R_{li} - R_{lo}$$

R_{lo}

Stefan-Boltzmann

law : $R_{lo} = (1 + \epsilon)$

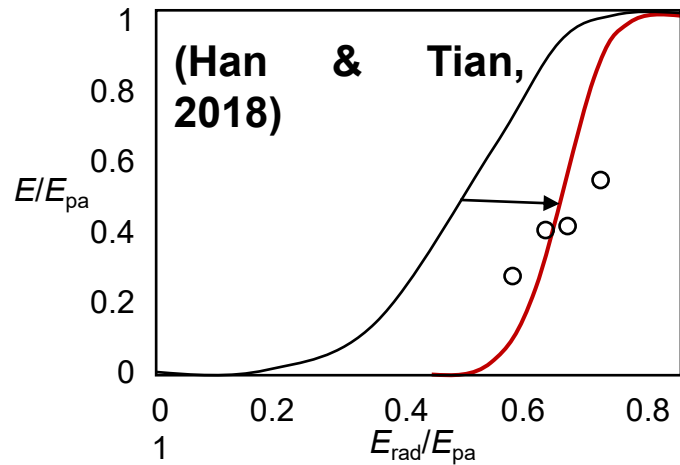


$$*R_{li} + \epsilon \sigma T_s^4$$

■ Problems of E_{po} estimation in the CR

- **Sigmoid CR (Han & Tian, 2018)**

$$\frac{E}{E_{Pen}} = \frac{1}{1 + m \left(\frac{E_{Pen}}{E_{rad}} - 1 \right)^n}$$



Problem:

- **Abandon the concept of Epa and Epo and has a vague physical interpretation**
- **Require calibration and prevents application over regions without prior knowledge**

■ Problems of E_{po} estimation in the CR



Temperature < Temperature

Solar radiation \approx Solar

radiation

Longwave radiation > Longwave

radiation

Humidity > Humidity

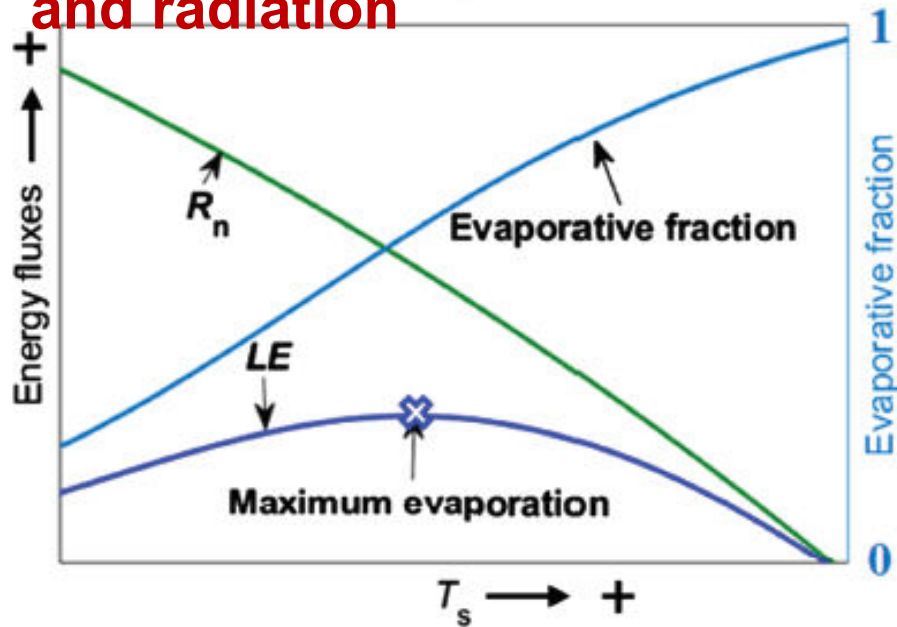
Wind speed \approx Wind speed

Key points:

- Unreasonable estimations of E_{po} hinder the development of CR
- How to estimate E_{po} following the definition (restore relevant variables to wet conditions) ?

■ The maximum evaporation theory (Yang & Roderick,

2019)
Acknowledging the inter-dependence between evaporation, surface temperature and radiation



Energy balance: $R_n - G = H + LE$

Bowen ratio: $\beta = H / LE$

E_p formulation: $LE = (R_n - G) * 1 / (1 + \beta)$

$T_s \uparrow$ $R_n \downarrow$ $1/(1+\beta)$ LE_{max}

LE_{max} naturally emerges from the coupling of

$R_n - T_s - E$

$$LE = \frac{R_{si} - R_{so} + \epsilon\sigma(T_s - \Delta T)^4 - \epsilon\sigma T_s^4}{1 + \beta(T_s)} = \frac{R_n(T_s)}{1 + \beta(T_s)}$$

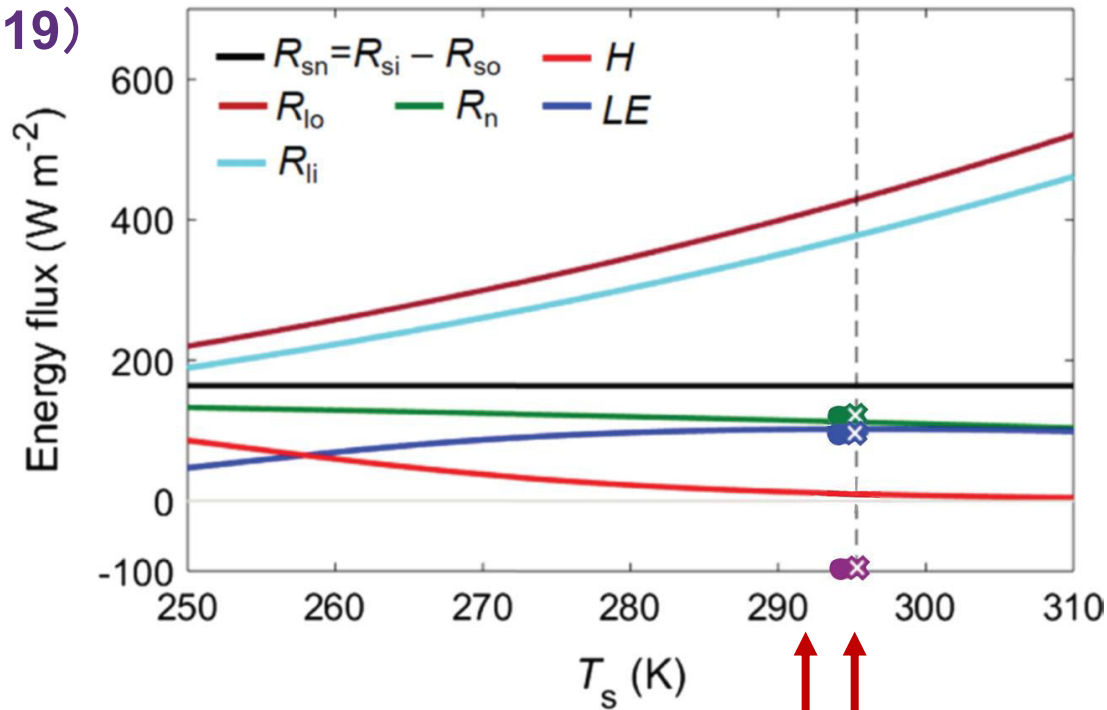
Energy balance : $R_n = R_{si} - R_{so} + R_{li} -$

R_{lo}

Stefan-Boltzmann law : $R_{lo} = (1 + \epsilon)$

■ The maximum evaporation theory (Yang & Roderick,

2019)



$$LE = \frac{R_{si} - R_{so} + \varepsilon\sigma(T_s - \Delta T)^4 - \varepsilon\sigma T_s^4}{1 + \beta(T_s)} = \frac{R_n(T_s)}{1 + \beta(T_s)}$$

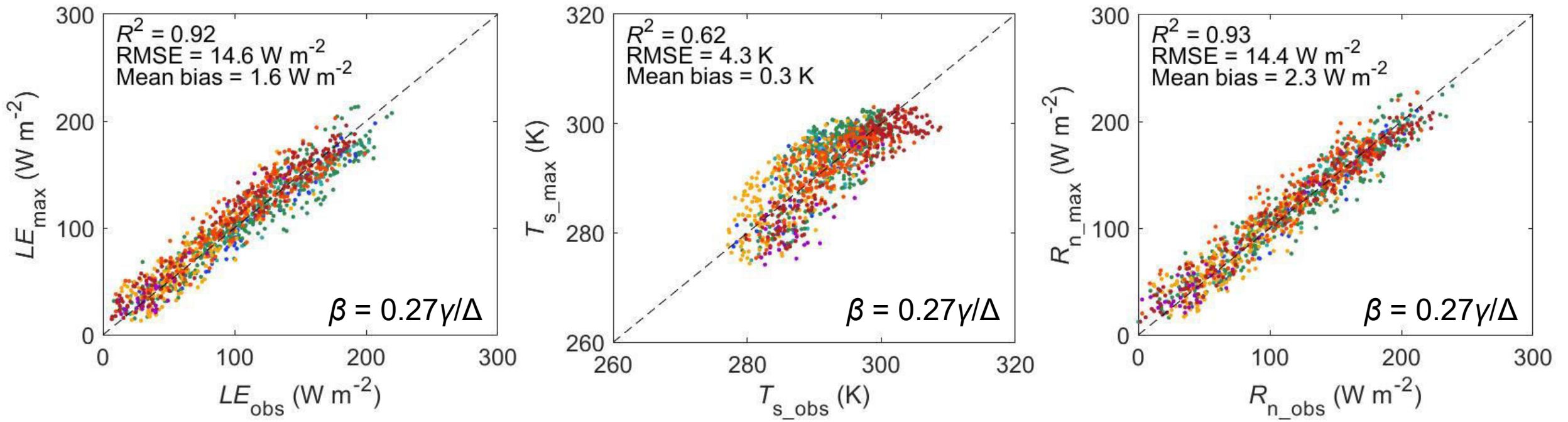
- This theory can restore E , R_n and T_s to water surface conditions.
- This theory has the potential to estimate E_p over land surfaces!

Observed LE	102.7 $W\ m^{-2}$
Observed T_s	293.8 K
Observed R_n	113.7 $W\ m^{-2}$

Max. LE	102.3 $W\ m^{-2}$
T_s at Max. LE	295.3 K
R_n at Max. LE	112.5 $W\ m^{-2}$

Consistent !

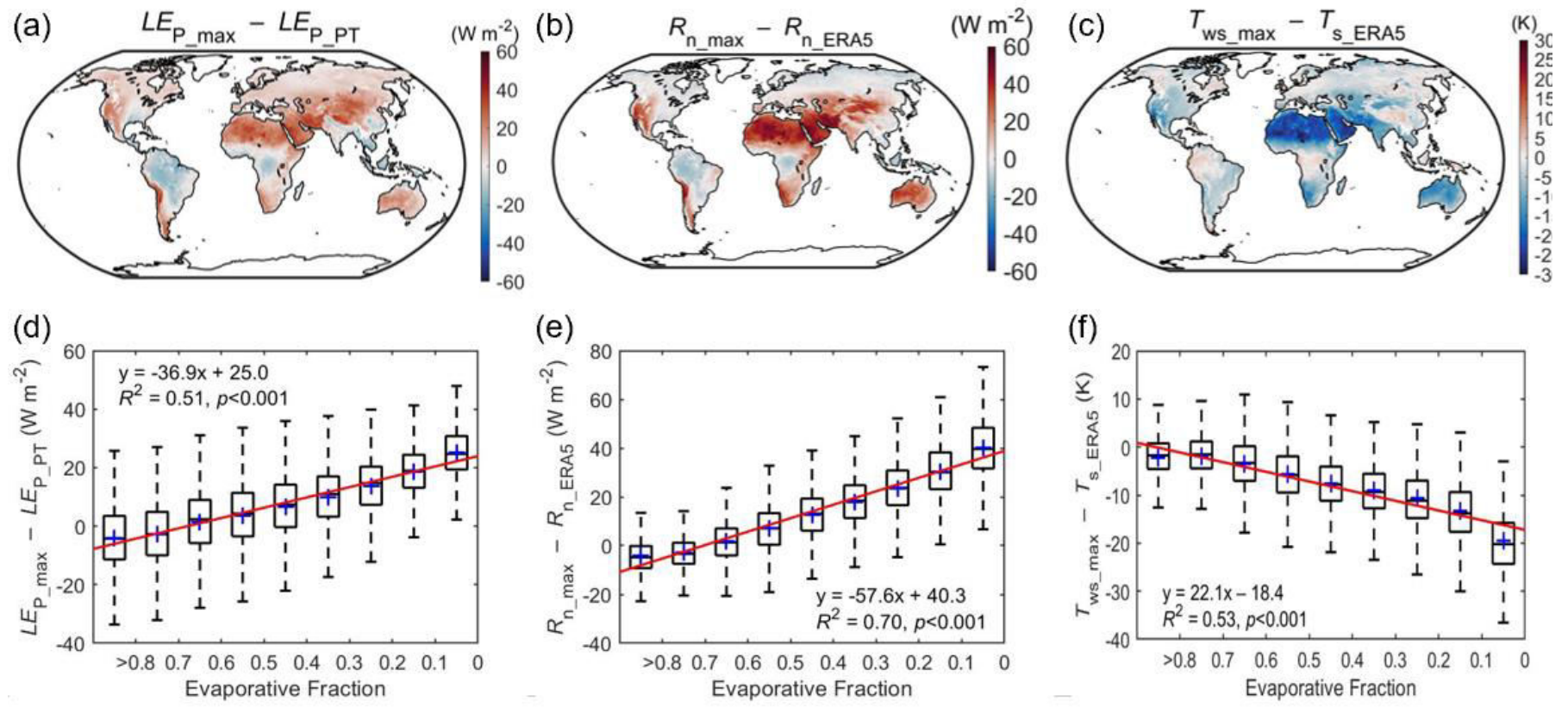
Validation of the maximum evaporation method over land surface (Tou et al., 2022, JGR)



The maximum evaporation method can restore E , T_s and R_n to wet conditions over land surface.

E_p estimates under both wet and dry conditions

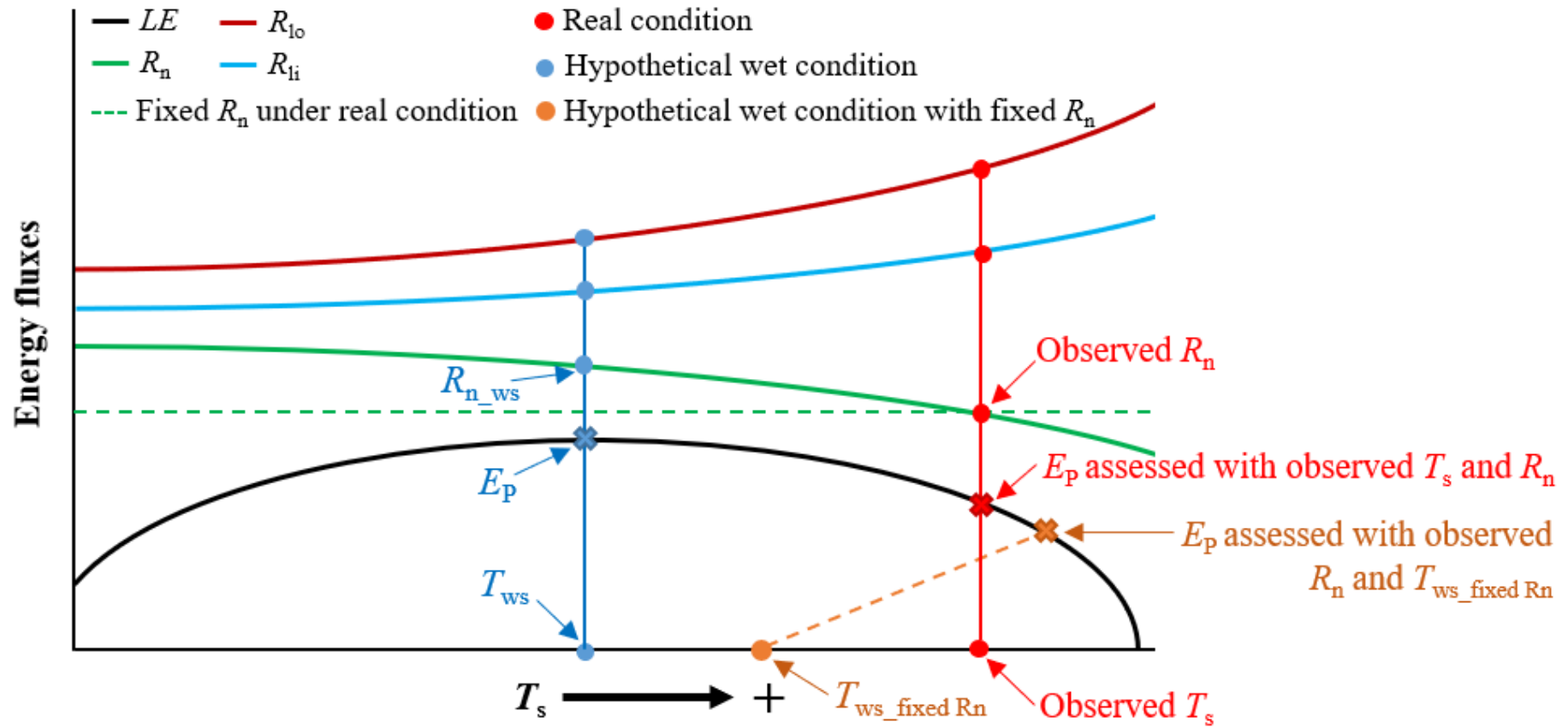
(Tu & Yang, 2022; WRR)



The underestimation of E_{p_PT} is caused the observed higher T_a and lower R_n under conditions.

E_p estimates under both wet and dry conditions

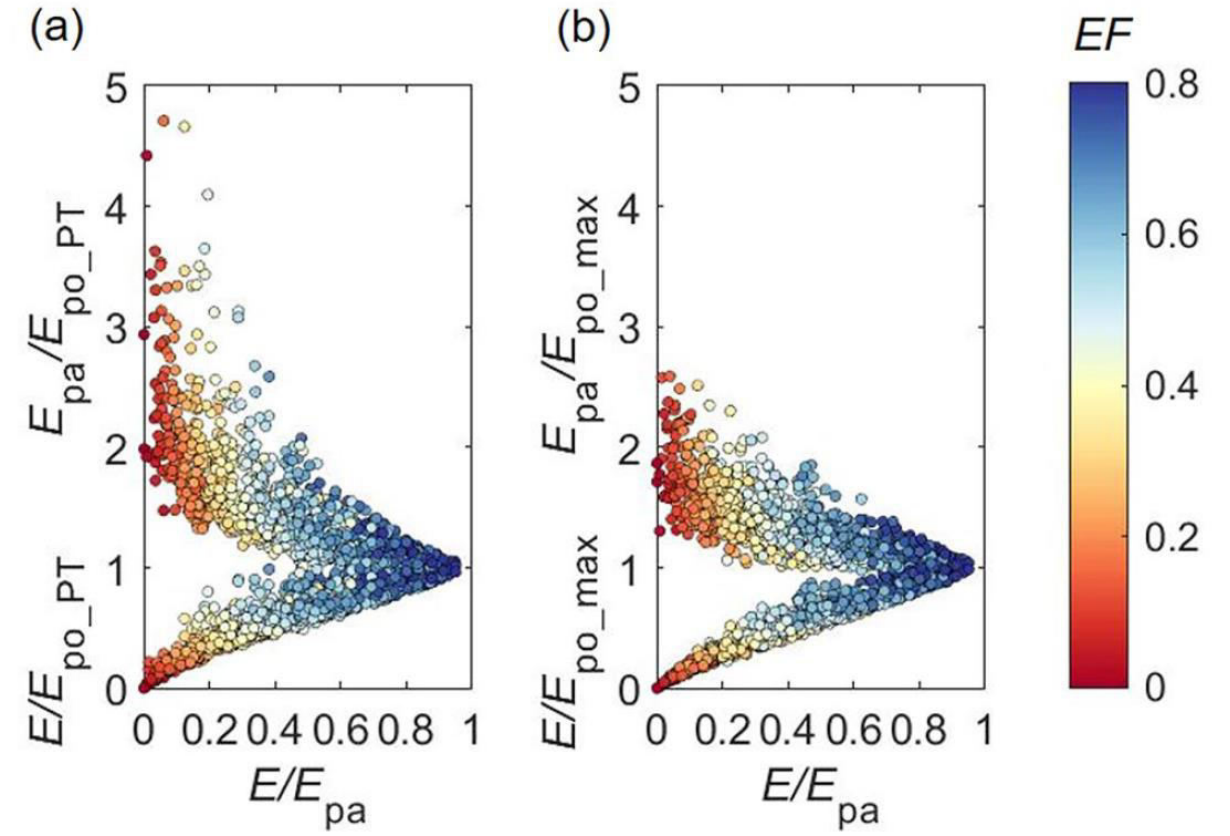
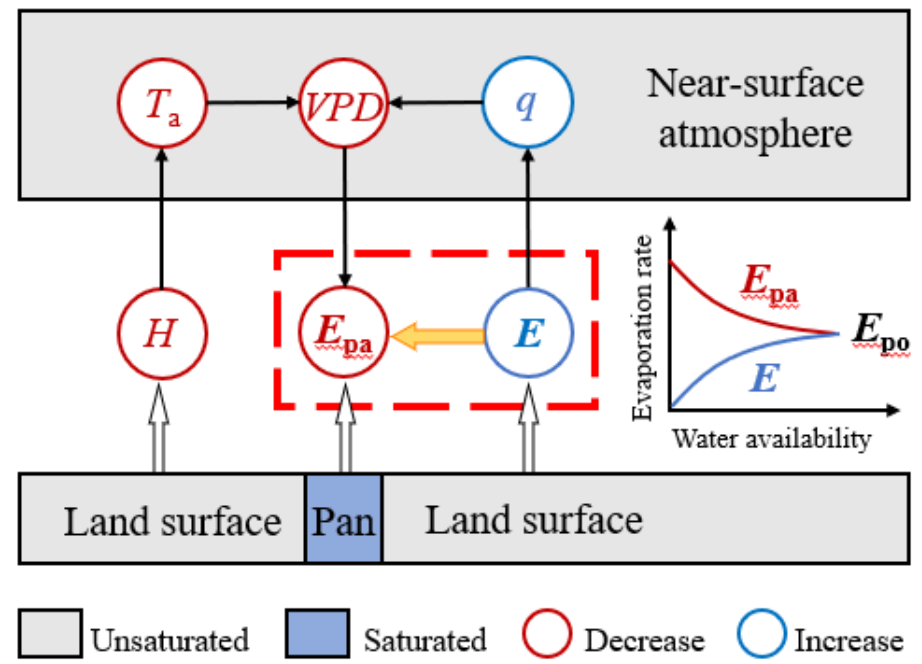
(Tu & Yang, 2022; WRR)



It is crucial to restore all the relevant variables (T_s, R_n) to wet conditions in estimating E_p !

CR with different E_{po} calculations

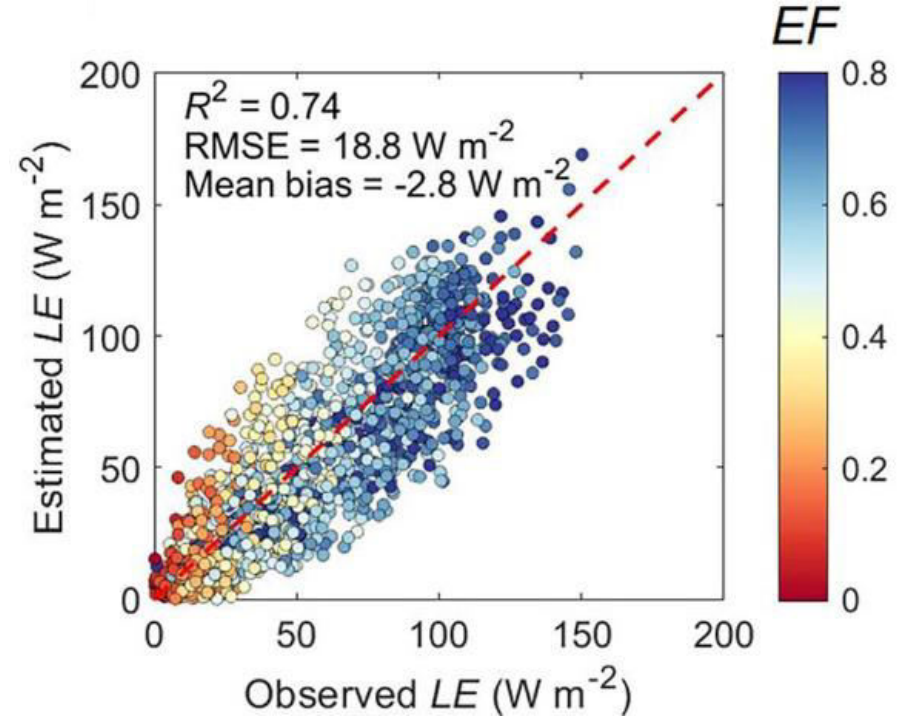
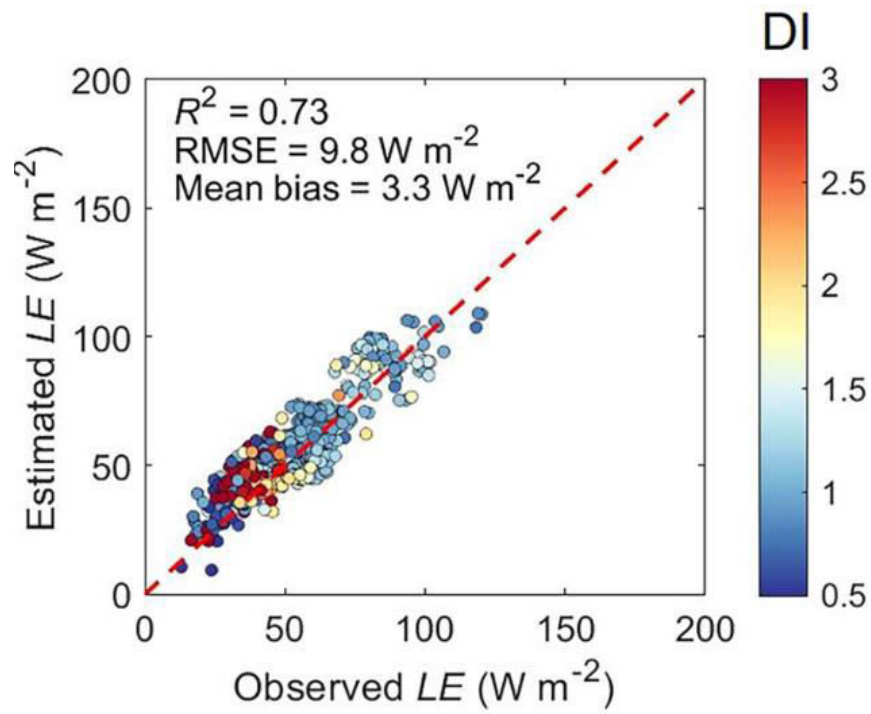
(Tu et al., 2023; WRR)



Adopting E_{po_max} considerably reduces the asymmetry of the complementarity in CR.

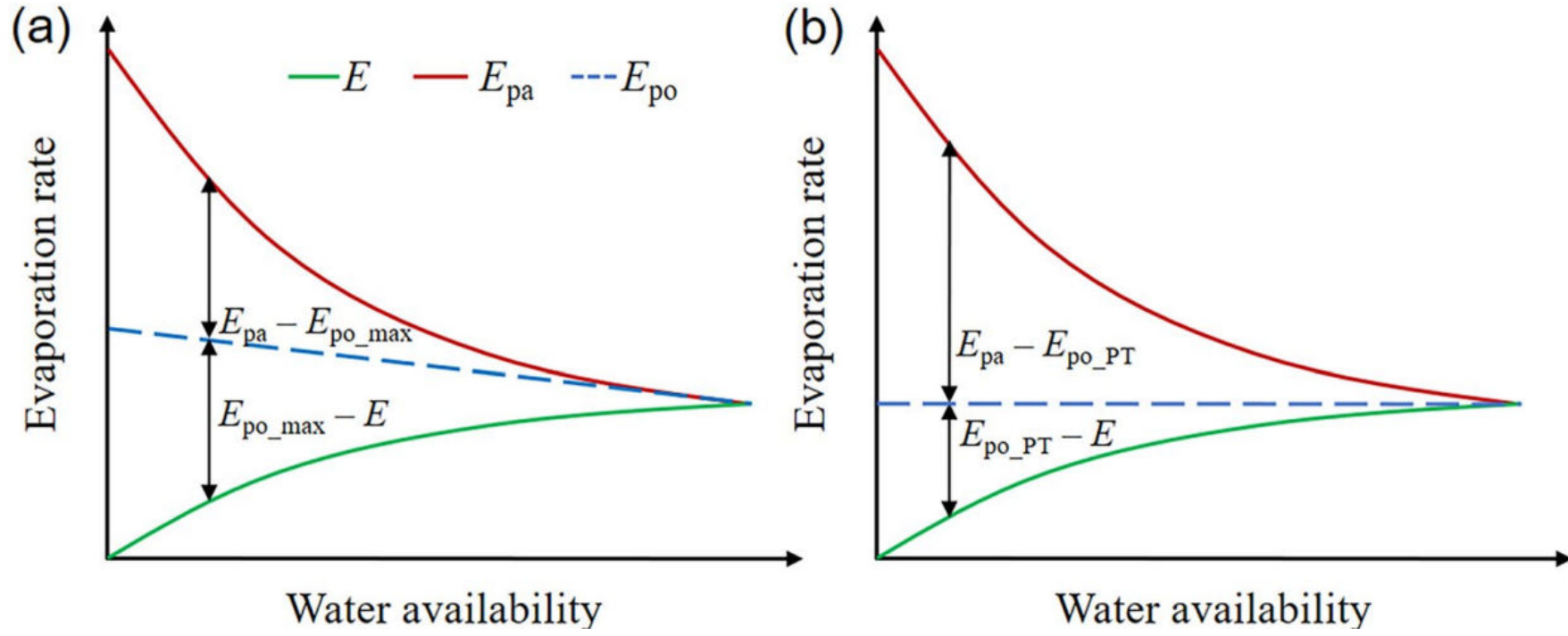
■ New CR model with E_{po_max}

(Tu et al., 2023; WRR)



A physically based CR model with E_{po_max} was validated to have good performance in E estimation.

Take-home messages



- The maximum evaporation method can **recover the variables to a hypothetical wet condition**
- Adopting E_{po_max} that conforms to the physical definition **reduces the asymmetry**

Thanks for your attention!

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