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The Flood Elasticity under Changing Conditions: An Application in a Tropical Monsoon Basin of Thailand

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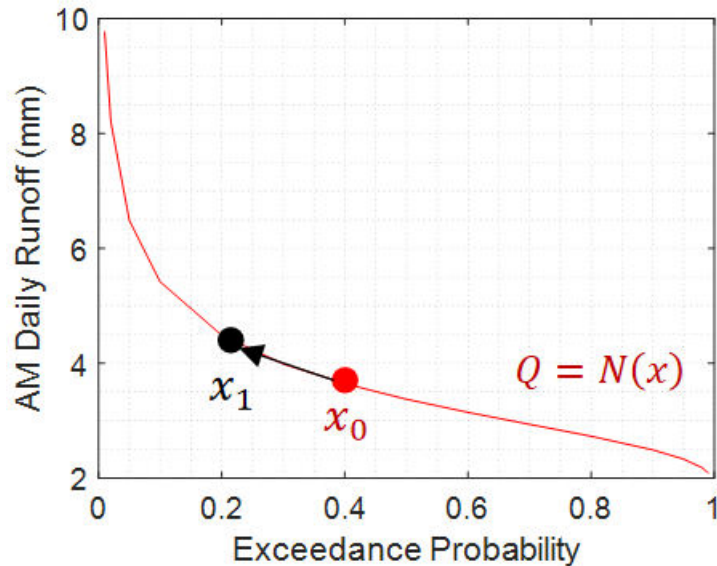
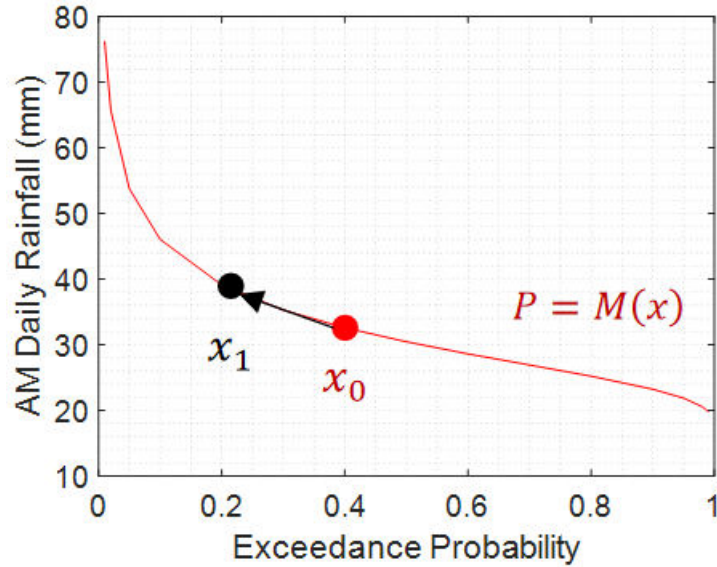
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- Definition: the **relative response** of **runoff** to the **relative change** of **precipitation**;
- Schaake et al. (1990): runoff elasticity estimates the relationship between rainfall and runoff **annually**;
- Breinl et al. (2021): storm-flood elasticity derived from **frequency curves** estimates the relationship between **daily** storm and flood peaks;
- Current **storm-flood elasticity analyses** normally regard the relationship as **constant** **regardless of changing conditions**.

- C-C Relationship: a 1°C increase in temperature leads to a resultant 7% increase in precipitation extremes;
- Intense studies have predicted **more and stronger precipitation extremes** across the world due to global warming (Donat et al., 2016, [NCC](#); Thackeray et al., 2022, [NCC](#));
- The **amplified flood magnitude** in the future is predicted in most parts of the world, while the contribution of climate change to flood amplification is the **strongest in Southeast Asia** (Winsemius et al., 2015, [NCC](#))

- Current elasticity analyses ignore changing conditions:
 - The frequency curves should tell changing stories.
 - ✓ Time-varying frequency curves via GAMLSS!
- Analyses on the future flood changes:
 - The changes of flood peaks should be quantified.
 - ✓ Time-varying flood frequency curves show!
 - The roles of different factors should be clarified.
 - ✓ Elasticity analyses on different factors!



Storm-Flood-Frequency-Equivalence:

- Suppose storm and the resultant flood are at the **same frequency**.
- Suppose exceedance probability decrease from x_0 to x_1 , design storm rise from P_0 to P_1 , design flood rise from Q_0 to Q_1 , correspondingly.

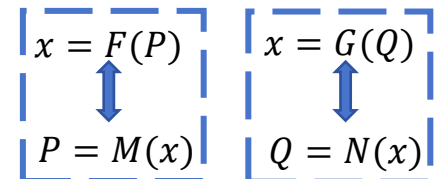
$$\begin{aligned} \varepsilon_0(x_0) &= \frac{\frac{Q_1 - Q_0}{Q_0}}{\frac{P_1 - P_0}{P_0}} \\ &= \frac{P_0 f(P_0)}{Q_0 g(Q_0)} \end{aligned}$$

$$\frac{\Delta Q}{Q} = \varepsilon_0 \frac{\Delta P}{P}$$

$$f(P) = \frac{\partial F}{\partial P} \quad g(Q) = \frac{\partial G}{\partial Q}$$

PDF Function

CDF Function



Inverse Function

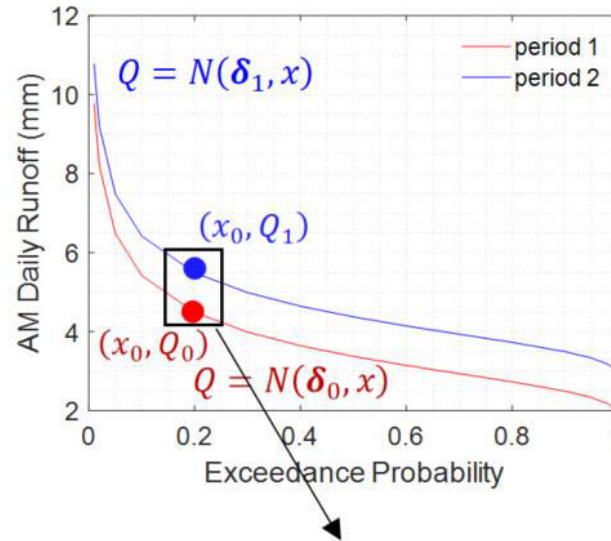
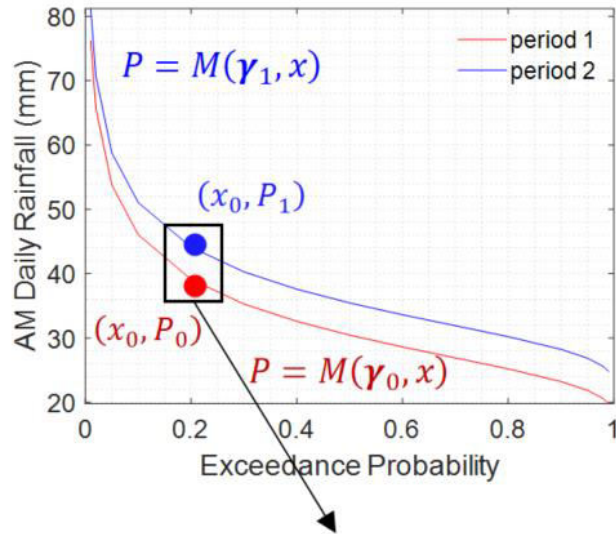
- Time-varying frequency curves:
 - **Covariates carrying information related to temporal variation** should be employed in the frequency analyses.
 - ✓ **Dew Point (DP)**: strongly positively correlated to precipitation extremes and flood peaks as indicated by C-C relationship (Ali et al., 2018; Worawiwat el al., 2021).
 - ✓ **Forest Cover (FC)**: strongly negatively correlated to flood peaks due to the hydrological restoration function of forest (Kim et al., 2019).

GAMLSS:

Regard distribution parameters as **time-varying**, herein the parameters change according to *DP* and *FC*.

$$P \sim GEV(\mu_{Pt}, \sigma_{Pt}, \xi_P)$$
$$\mu_{Pt} = \alpha_{P0} + \alpha_{P1}DP_t$$
$$\sigma_{Pt} = \beta_{P0} + \beta_{P1}DP_t$$

$$Q \sim GEV(\mu_{Qt}, \sigma_{Qt}, \xi_Q)$$
$$\mu_{Qt} = \alpha_{Q0} + \alpha_{Q1}DP_t + \alpha_{Q2}FC_t$$
$$\sigma_{Qt} = \beta_{Q0} + \beta_{Q1}DP_t + \beta_{Q2}FC_t$$

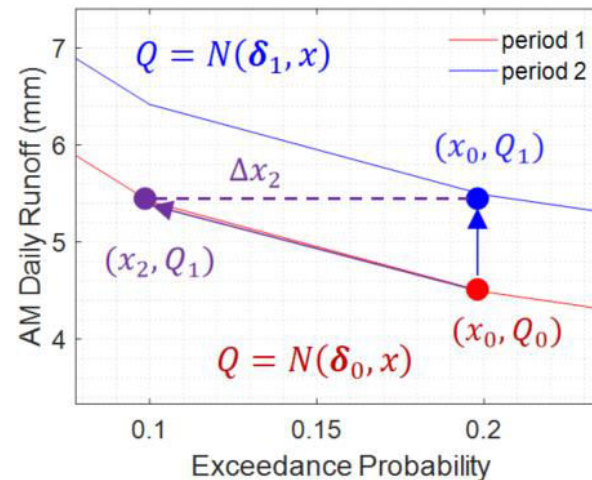
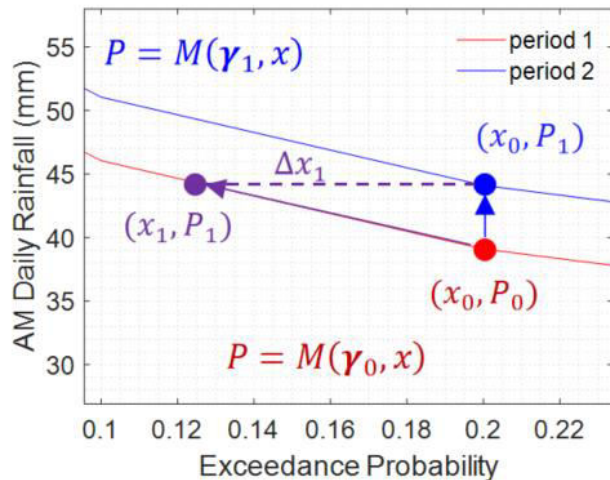


CDF Function

$$\begin{array}{|c|c|}
 \hline
 x = F(\boldsymbol{\gamma}, P) & x = G(\boldsymbol{\delta}, Q) \\
 \hline
 \updownarrow & \updownarrow \\
 P = M(\boldsymbol{\gamma}, x) & Q = N(\boldsymbol{\delta}, x) \\
 \hline
 \end{array}$$

Inverse Function

$$\boldsymbol{\gamma} = (DP) \quad \boldsymbol{\delta} = (DP, FC)$$



$$\varepsilon_C = \frac{\frac{\partial G}{\partial DP}}{\frac{\partial F}{\partial DP}} \varepsilon_0$$

$$\varepsilon_F = \frac{\frac{\partial G}{\partial DP} FC_0}{\frac{\partial F}{\partial DP} P_0} \frac{\partial M}{\partial DP} \varepsilon_0$$

As the parameters changes from $\boldsymbol{\gamma}_0$ to $\boldsymbol{\gamma}_1$ ($\boldsymbol{\delta}_0$ to $\boldsymbol{\delta}_1$), the frequency curves shift from the red one to the blue one.

$$\frac{\Delta Q}{Q} = \varepsilon_C \frac{\Delta P}{P} + \varepsilon_F \frac{\Delta FC}{FC}$$

Study Region: UCPRB

3.CASE STUDY

The Upper Chao Phraya River
Basin (UCPRB)



Area: 110,553 km²

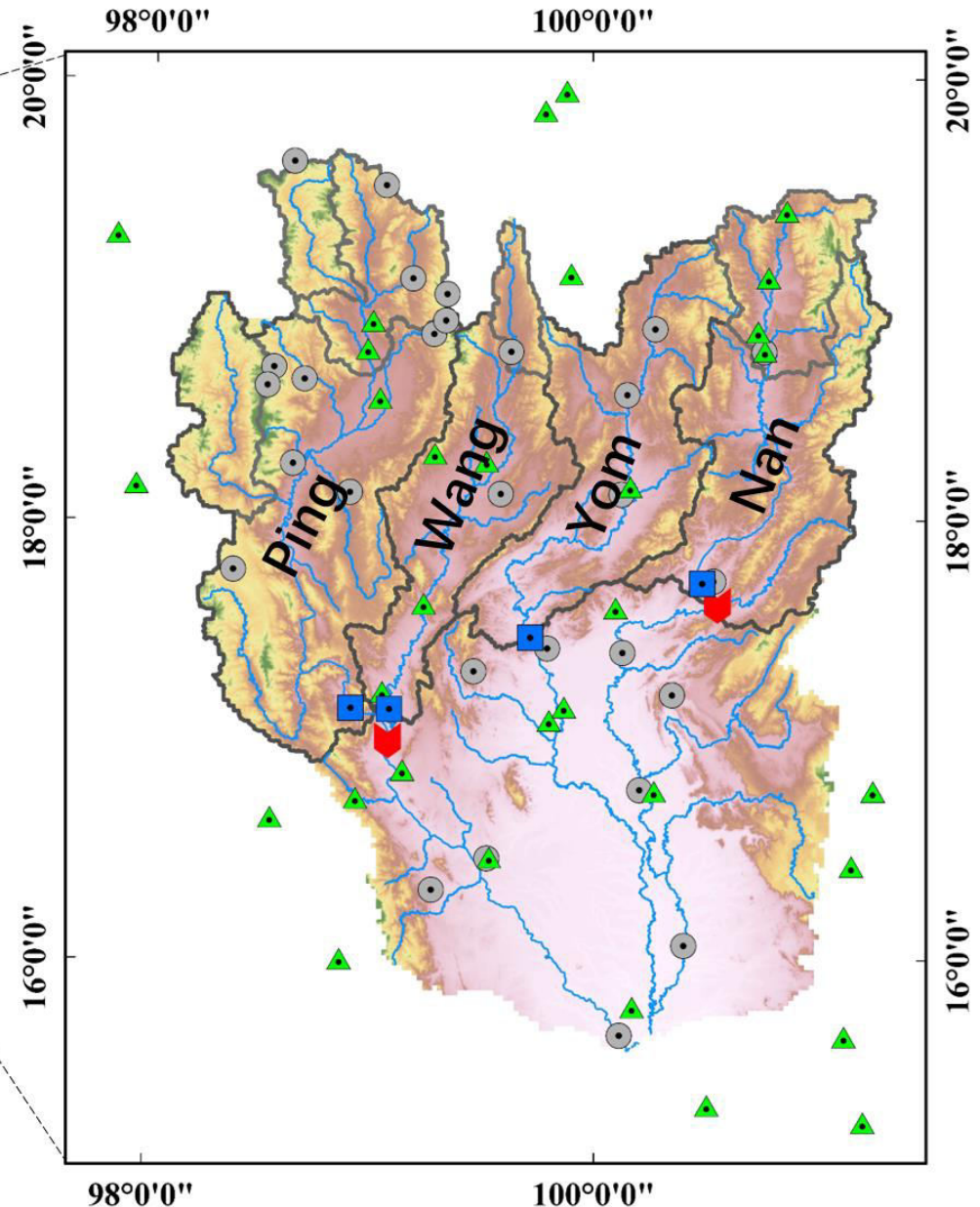
Annual Pre: 1100 mm

Legend

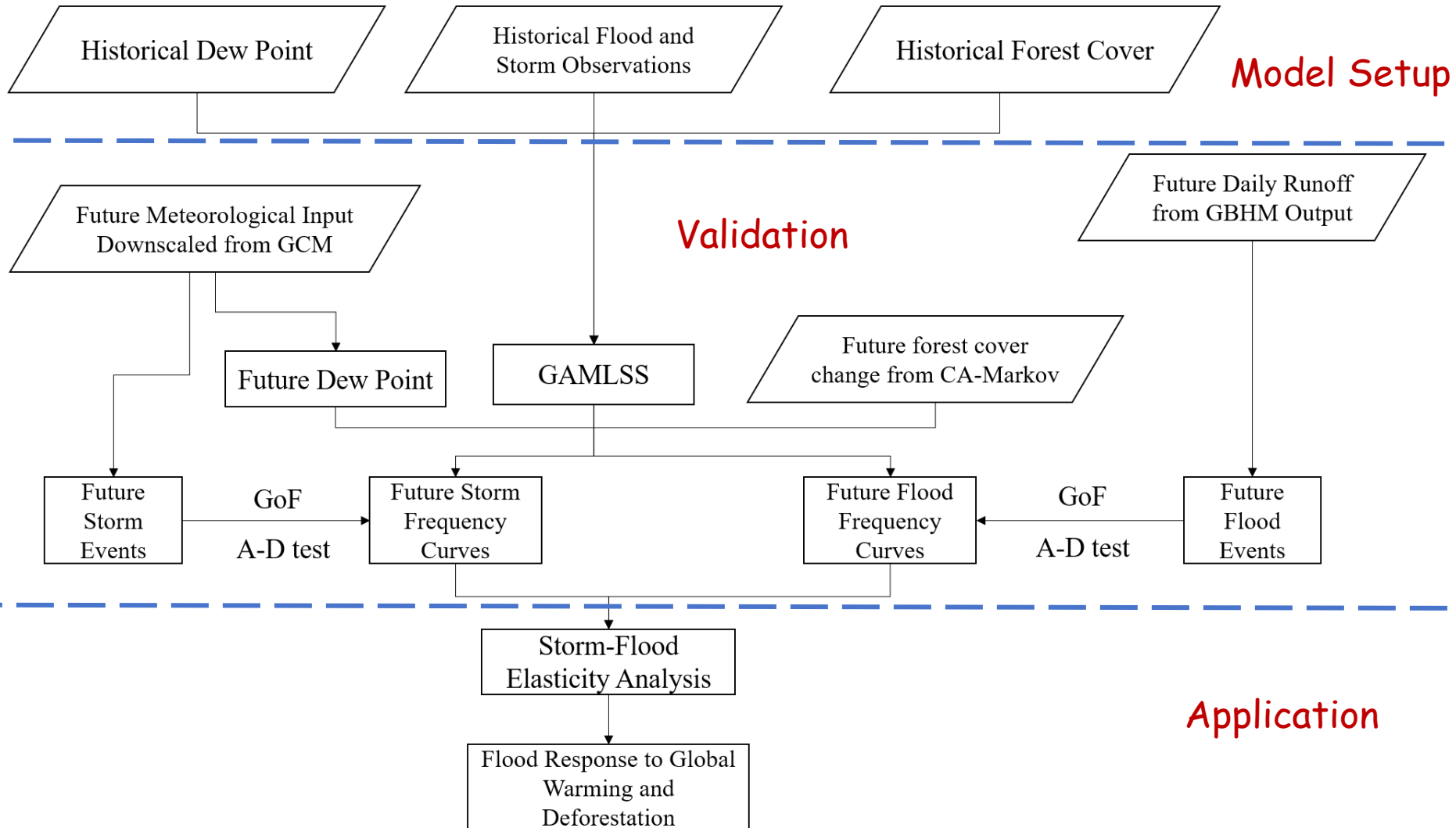
Elevation/m
Value

High : 2559 Low : 1

-  Meteorological Stations
-  Rain Gauges
-  Hydrological Stations
-  Reservoirs
-  River_Network



Variable	Temporal resolution	Spatial resolution	Time period available	Source
Runoff	Daily	/	1981-2014	Royal Irrigation Department
historical values: mainly observations		/	2015-2099	Yang et al. (2023)
used for model setup				
Precipitation	Daily	/	1981-2014	Thailand Meteorological Department
future values: from CMIP6 and hydrological models (Yang et al. 2023)		2.5 km	2015-2099	Yang et al. (2023)
Relative Humidity	Daily	/	1981-2014	Thailand Meteorological Department
used for model validation and prediction		2.5 km	2015-2099	Yang et al. (2023)
Mean Temperature	Daily	/	1981-2014	Thailand Meteorological Department
		2.5 km	2015-2099	Yang et al. (2023)
DEM	/	90 m	/	HydroSHEDS
Land use	Annual	100 m	1981-2014	Zhao et al. (2022)
			2015-2099	Yang et al. (2023)



Storm and flood events are captured using the **Annual Maximum** method.

The parameters in the GAMLSS model are then estimated using **MCMC**.

- The GoF is quantified using **Anderson-Darling test**.
- The p-values are all larger than 0.05, so the models are accepted.

$$P \sim GEV(\mu_{Pt}, \sigma_{Pt}, \xi_P)$$

$$\begin{aligned} \mu_{Pt} &= \alpha_{P0} + \alpha_{P1}DP_t \\ \sigma_{Pt} &= \beta_{P0} + \beta_{P1}DP_t \end{aligned}$$

$$Q \sim GEV(\mu_{Qt}, \sigma_{Qt}, \xi_Q)$$

$$\begin{aligned} \mu_{Qt} &= \alpha_{Q0} + \alpha_{Q1}DP_t + \alpha_{Q2}FC_t \\ \sigma_{Qt} &= \beta_{Q0} + \beta_{Q1}DP_t + \beta_{Q2}FC_t \end{aligned}$$

	α_{P0}	α_{P1}	β_{P0}	β_{P1}	p-value
Ping	28.99	3.24	7.60	2.91	0.967
Wang	27.95	5.50	6.24	1.72	0.686
Yom	28.73	7.98	10.51	2.71	0.738
Nan	26.06	9.49	10.74	3.71	0.742

parameters in storm distribution

	α_{Q0}	α_{Q1}	α_{Q2}	β_{Q0}	β_{Q1}	β_{Q2}	p-value
Ping	2.99	2.19	-0.48	1.41	0.10	-0.11	0.702
Wang	2.31	2.14	-0.25	1.80	0.05	-0.43	0.996
Yom	3.02	1.99	-0.40	2.38	0.26	-0.58	0.970
Nan	6.97	3.40	-1.00	4.64	0.50	-0.69	0.810

parameters in flood distribution

- Future distributions can then be estimated using the *DP* and *FC* predictions.

- The future is divided into 3 periods accordingly:

✓ Near-Future (2026-2055)

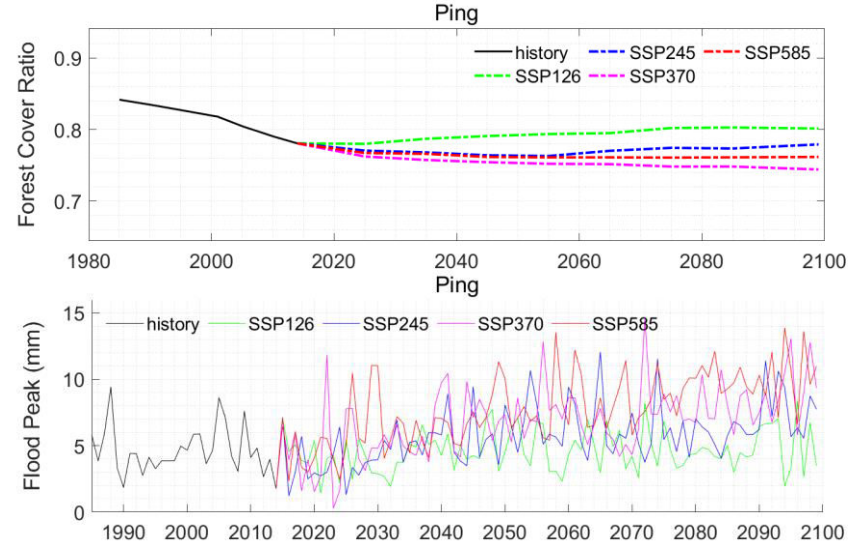
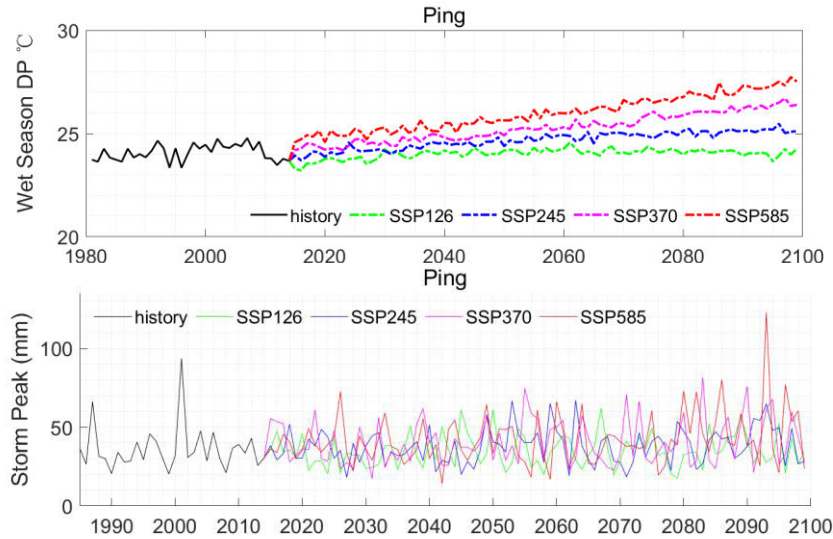
✓ Mid-Future (2046-2075)

✓ Far-Future (2070-2099)

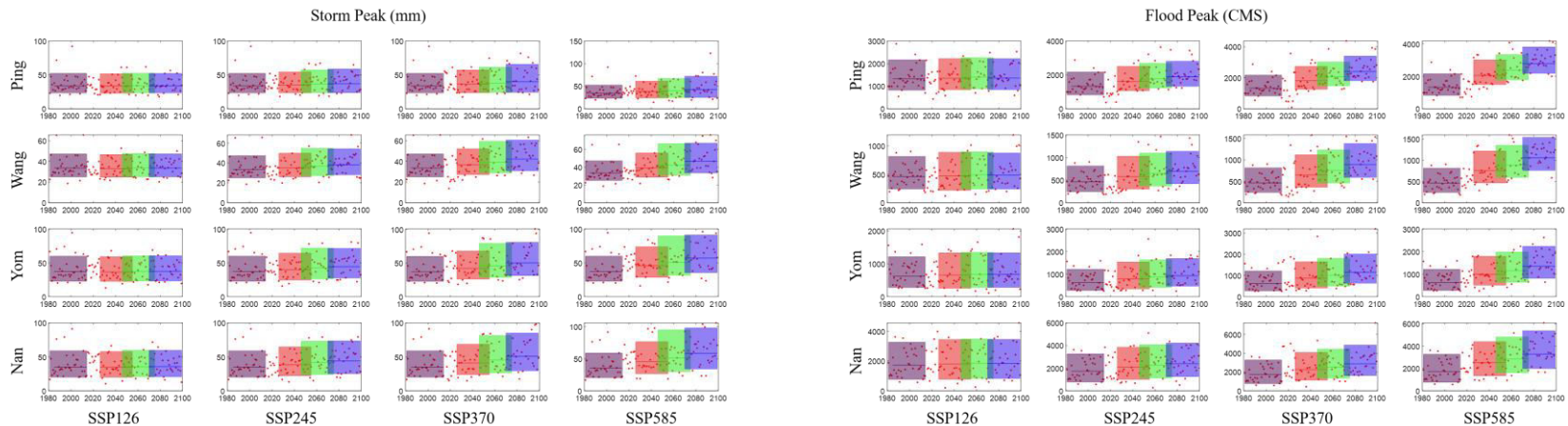
Result: Model Performance

3.CASE STUDY

The **covariates** (DP , FC), **storm** and **flood** peaks from **process-based models** (PBM) in the Ping subbasin.

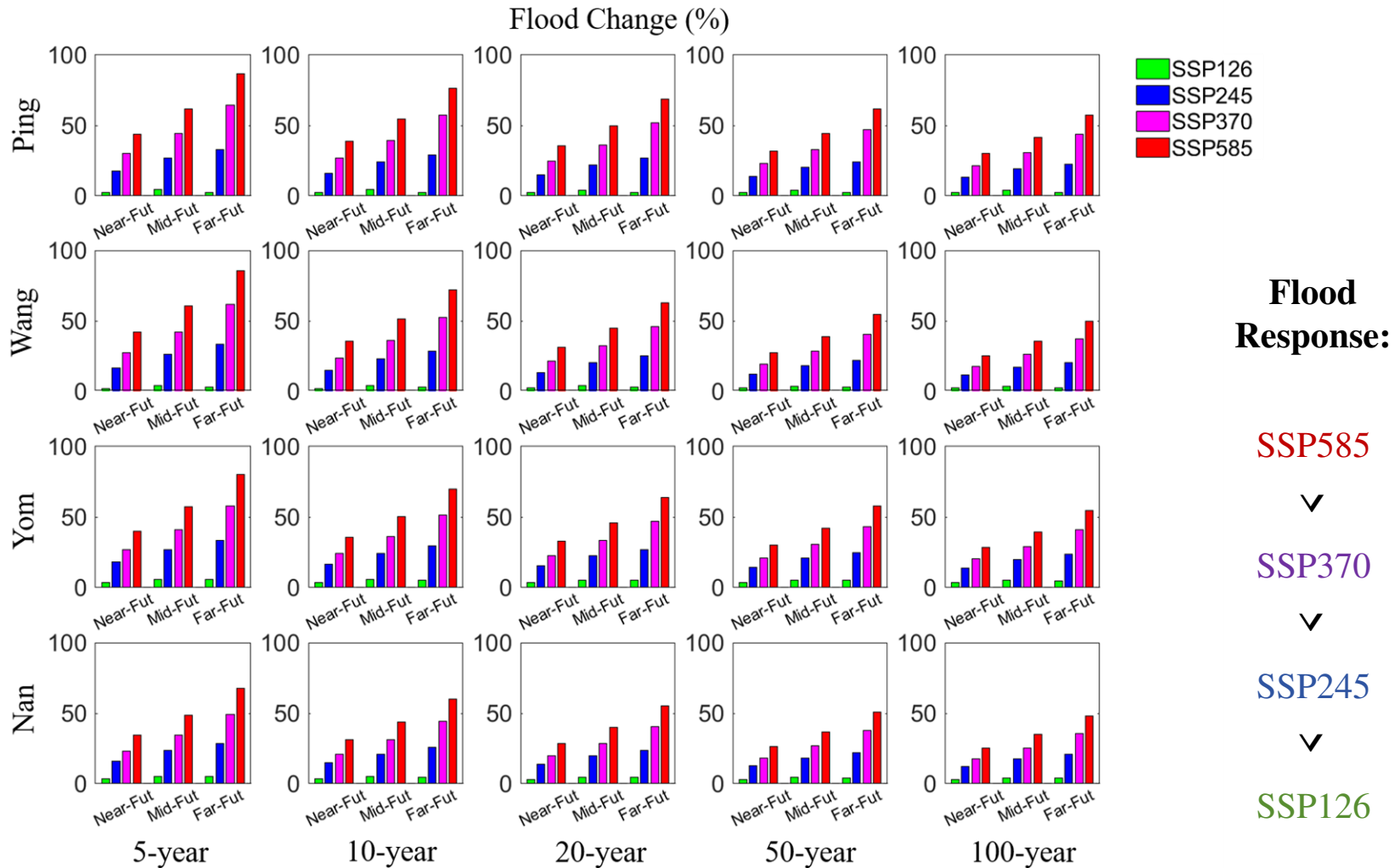


The future storm and flood distributions from **PBM** and **GAMLSS** are compared using **A-D test**, and the result shows that GAMLSS captures the distributions well.



Result: Future Changes in Floods

3.CASE STUDY



- The flood peaks keep increasing till the late 21st century.
- Under the SSP585 scenario where the flood increase is the strongest, the flood increase at late 21st century can reach 47%~58%.

Flood Elasticity

- **Positive** correlation with storm and **negative** correlation with forest;
- Different storm-flood elasticity under stationary and nonstationary cases;
- Weaker elasticity for larger events.

Contribution Analyses

Climate-change-induced

flood change

$$\frac{\Delta Q}{Q} = \left[\varepsilon_C \frac{\Delta P}{P} \right] + \left[\varepsilon_F \frac{\Delta FC}{FC} \right]$$

Deforestation-induced
flood change

$$\varepsilon_0 = \frac{P_0 \frac{\partial F}{\partial P}}{Q_0 \frac{\partial G}{\partial Q}}$$

$$\varepsilon_C = \frac{\frac{\partial G}{\partial DP}}{\frac{\partial F}{\partial DP}} \varepsilon_0$$

$$\varepsilon_F = \frac{\frac{\partial G}{\partial DP} FC_0}{\frac{\partial F}{\partial DP} P_0} \frac{\partial M}{\partial DP} \varepsilon_0$$

ε_0	return period				
	5	10	20	50	100
Ping	1.11	1.08	1.05	1.03	1.01
Wang	1.67	1.53	1.43	1.33	1.28
Yom	1.35	1.26	1.20	1.15	1.11
Nan	1.20	1.14	1.11	1.07	1.05

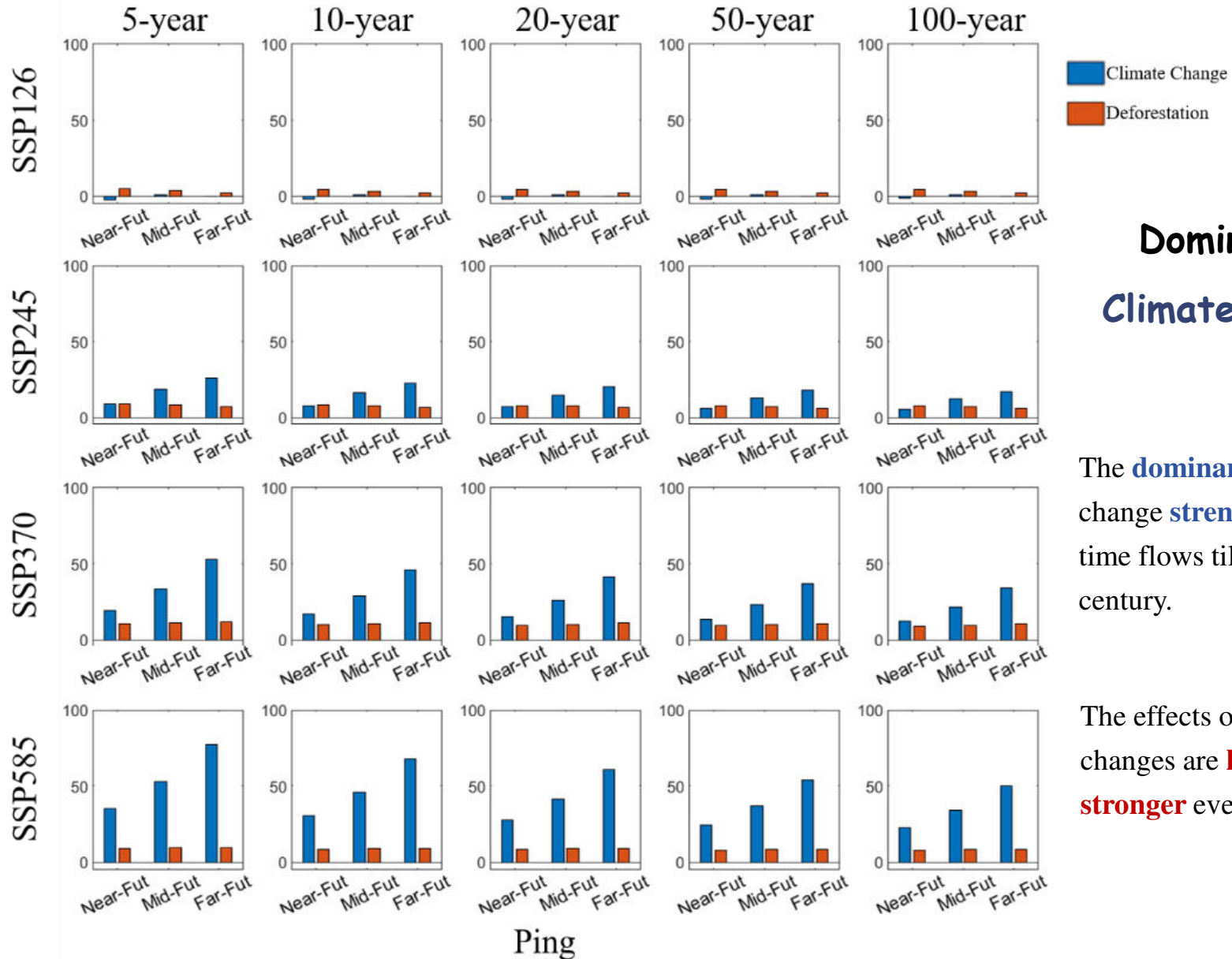
ε_C	return period				
	5	10	20	50	100
Ping	2.27	1.78	1.48	1.23	1.09
Wang	1.93	1.56	1.32	1.10	0.98
Yom	1.36	1.19	1.08	0.98	0.92
Nan	0.90	0.80	0.73	0.67	0.63

ε_F	return period				
	5	10	20	50	100
Ping	-1.43	-1.37	-1.32	-1.27	-1.25
Wang	-0.62	-0.61	-0.60	-0.59	-0.58
Yom	-0.71	-0.69	-0.68	-0.67	-0.66
Nan	-0.63	-0.60	-0.57	-0.55	-0.54

Result: Contribution Analyses

3.CASE STUDY

Flood Peak Response to Changing Conditions (%)



Dominance:
Climate Change

The **dominance** of climate change **strengthens** as time flows till late 21st century.

The effects of **land use** changes are **larger** for **stronger** events.

- Flood elasticity analyses under changing conditions:
 - ✓ **Employment of covariates:** Enabling the predictions of future floods;
 - ✓ **Statistical model (GAMLSS):** Validated by process-based models;
 - ✓ **Flood elasticity:** Weaker for stronger events;
 - ✓ **Storm-flood elasticity:** Different values under the nonstationary case.
- Future flood peak changes in the UCPRB:
 - ✓ **Future flood:** Increasing till late 21st century;
 - ✓ **Scenario:** SSP585>SSP370>SSP245>SSP126;
 - ✓ **Climate change:** Increasing dominance as time flows;
 - ✓ **Deforestation:** Larger effects for stronger events.

Zhao.B., T. Wang, D. Yang, S. Yang, W. Lu and J. Santisirisomboon (2023), The impacts of climatic and land surface characteristics on the storm-flood relationship in a tropical monsoon basin of Thailand, *Journal of Hydrology*, 616, doi.org/10.1016/j.jhydrol.2022.128809

Zhao.B., T. Wang, D. Yang, S. Yang and J. Santisirisomboon, The flood elasticity under changing conditions: an application in a tropical monsoon basin of Thailand, *Weather and Climate Extremes* (Under Review)



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Thanks for Your Attention!

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